

Sensitivity of stand transpiration to wind velocity in a mixed broadleaved deciduous forest



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ABSTRACT

Wind velocity (U) within and above forest canopies can alter the coupling between the vapor-saturated sub-stomatal airspace and the drier atmosphere aloft, thereby influencing transpiration rates. In practice, however, the actual increase in transpiration with increasing U depends on the aerodynamic resistance (R_A) to vapor transfer compared to canopy resistance to water vapor flux out of leaves (R_C , dominated by stomatal resistance, R_{stom}), and the rate at which R_A decreases with increasing U . We investigated the effect of U on transpiration at the canopy scale using filtered meteorological data and sap flux measurements gathered from six diverse species of a mature broadleaved deciduous forest. Only under high light conditions, stand transpiration (E_C) increased slightly (6.5%) with increasing U ranging from ~ 0.7 to ~ 4.7 m s⁻¹. Under other conditions, sap flux density (J_s) and E_C responded weakly or did not change with U . R_A , estimated from Monin–Obukhov similarity theory, decreased with increasing U , but this decline was offset by increasing R_C , estimated from a rearranged Penman–Monteith equation, due to a concurrent increase in vapor pressure deficit (D). The increase of R_C with D over the observed range of U was consistent with increased R_{stom} by $\sim 40\%$ based on hydraulic theory. Except for very rare half-hourly values, the proportion of R_A to total resistance (R_T) remained $<15\%$ over the observed range of conditions. These results suggest that in similar forests and conditions, the direct effect of U reducing R_A and thus increasing transpiration is negligible. However, the observed U – D relationship and its effect on R_{stom} must be considered when modeling canopy photosynthesis.

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1. Introduction

Because of their expanse, forests greatly impact the exchange of mass and energy between the biosphere and atmosphere with consequences ranging from effects on local weather conditions (Juang et al., 2007; Konings et al., 2011) to influences on global atmospheric circulation (Avisar and Werth, 2005; Bauerle and Bowden, 2011). Forests also play an important role in the provision of ecosystem services, including water yield, biomass production and carbon sequestration. Improved understanding of these exchanges and benefits involves a combination of observations and models. Estimates of mass (including carbon and water) and energy exchange between the biosphere and atmosphere commonly rely on modeled

and measurement-based estimates of a series of resistances to these exchanges, each affected by biological and physical factors.

The flux of water vapor through stomata (i.e. transpiration) is an important flux since it represents the major component of water returned to the atmosphere from forests. Transpiration is also tightly coupled to CO₂ exchange (i.e. canopy photosynthesis or gross primary production) because stomata serve as the common port for the fluxes of these gases. Thus, tree and ecosystem physiologists often investigate the sources of variation of resistances to water vapor flux in lieu of the variation of resistances to CO₂ flux (Kim et al., 2008; Schäfer et al., 2003), and models often computationally link canopy CO₂ uptake to transpiration via the carbon-based water-use efficiency (Running and Hunt, 1993). Transpiration is convenient flux to use in such analysis because, unlike CO₂ flux, the dynamics of its exchange is straightforward to measure (Granier, 1987; Oishi et al., 2008; Oren et al., 1998; Schäfer et al., 2000).

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Table 1

Field studies that have investigated relationships between transpiration (or sap flux) and wind velocity.

Correlation	Site description (species)	Wind velocity (m s ⁻¹)	Tree stage (age in years)	LAI ^a	SLA ^b (cm ² g ⁻¹)	Source
Positive	Hardwood, South Australia (<i>Eucalyptus globulus</i>)	–	Young (7)	–	62–192 ^d	Taylor et al. (2001)
	Boreal mixedwood, Alberta, Canada (<i>Picea glauca</i> , <i>P. balsamifera</i> , <i>Betula papyrifera</i>)	0–2	Mature (45–50)	–	–	Bladon et al. (2006)
	Shrub, Tengger desert, China (<i>Artemisia ordosica</i>)	–	(10)	0.5–0.9	–	Lei et al. (2010)
	Temperate hardwood, Estonia (<i>B. pendula</i> , <i>P. tremula</i> + <i>P. tremuloides</i> ^c)	0–4	Seedling (1)	–	–	Kupper et al. (2011)
	Temperate hardwood, China (<i>Populus alba</i>)	0–0.6	Young (7–10)	0–2	100–120 ^d	Xu et al. (2006)
Negative	Shrub, island of Kauai, USA (<i>Coffea arabica</i> L.)	1–3	(1–5.3)	0.7–6.7	–	Gutiérrez et al. (1994)
	Hardwood, Wisconsin, USA (<i>P. nigra</i> + <i>P. maximowiczii</i> ^c)	0–4.5	Young	–	–	Zalesny et al. (2006)
Weak or none	Hardwood, Plauzat, France (<i>Juglans regia</i>)	0–4	Mature (20)	9	82–398 ^d	Daudet et al. (1999)
	Softwood, Chiba, Japan (<i>Cryptomeria japonica</i>)	0–1.5	Young (8)	3.7	–	Komatsu et al. (2006)
	Crop, Nebraska, USA (<i>Zea mays</i> L.)	0–7	–	1–5	–	Irmak and Mutiibwa (2010)
	Temperate hardwood, Durham, USA (<i>Liriodendron tulipifera</i> , <i>Liquidambar styraciflua</i> , <i>Carya tomentosa</i> , <i>Quercus alba</i> , <i>Q. michauxii</i> , <i>Q. phellos</i>)	0.8–4.7	Mature (80–100)	7.03	102–174	This study

^a Leaf area index.^b Specific leaf area.^c Hybrid between the two species.^d General values from various studies.

Resistance to transpiration (R_T) can be partitioned into two serial components, bulk surface resistance (R_{sur}) and aerodynamic resistance (R_A),

$$R_T = R_{sur} + R_A \quad (1)$$

both of which depend on leaf physiology, canopy architecture and the atmospheric conditions above and within the canopy (Jones, 1992; Kumagai et al., 2004). R_{sur} is generally attributed to canopy resistance (R_C) and soil (or forest floor) resistance (R_{soil}) acting in parallel (Allen, 2005).

$$R_{sur} = \frac{R_C R_{soil}}{R_C + R_{soil}} \quad (2)$$

Soil evaporation is minor under high canopy leaf area, making R_C effectively equivalent to R_{sur} when leaf area index (LAI) is greater than 3 (Kelliher et al., 1995; Waring, 2007). R_C can be expressed as a combination of stomatal resistance (R_{stom}) and leaf boundary layer resistance (R_{bl}) in series (Aphalo and Jarvis, 1993):

$$R_C = R_{stom} + R_{bl} \quad (3)$$

R_{stom} reflects complex physiological responses of stomata to variation in atmospheric conditions, such as light availability, air temperature (T_a), humidity, CO₂ concentration (Jarvis, 1976). R_{bl} is the resistance to vapor transfer within the air layer near leaf surface, and can be affected by wind velocity (U), leaf shape, leaf size, edge characteristics, and surface structure (Schuepp, 1993; Stokes et al., 2006). R_{bl} is generally much smaller than R_{stom} unless leaf size is large or U is low (Aphalo and Jarvis, 1993; Schuepp, 1993).

Over time scales in which soil moisture and plant hydraulic properties (e.g. soil-to-leaf path-length, sapwood-to-leaf area ratio, hydraulic conductivity) can be considered invariable, R_C is largely determined by R_{stom} . Jarvis (1976) described stomatal conductance (the inverse of R_{stom}) as a form of multiplicative function with

photon flux density, vapor pressure deficit (D), T_a , ambient CO₂ concentration, and leaf water potential. Many formulations for R_{stom} , ranging from empirical to theoretically-based functions have been proposed and used (Holmgren et al., 1996; Jonard et al., 2011; Launiainen et al., 2011; Pataki and Oren, 2003; Stewart, 1988), but except for allowing interactions among the controlling variables (Ward et al., 2008), this approach remained fundamentally unaltered. The other primary component of R_T , R_A , is determined by the physical conditions above and within the canopy, such as U , atmospheric stability, and canopy structure (Kumagai et al., 2004; Lindroth, 1993; Morris et al., 1998). Although estimates of R_A in forests are usually small relative to R_C , reflected in that R_A is neglected in some biosphere–atmosphere exchange models (Granier et al., 2000; Kelliher et al., 1993; Köstner et al., 1992), few studies have actually quantified the effect of U on R_A and, in-turn, on transpiration (Table 1). The results of these studies are inconclusive: some studies observed increased transpiration with U (Chu et al., 2009; Grace and Russel, 1982; Kitaya et al., 2004); others observed a decline (Caldwell, 1970; Campbell-Clouse, 1998; Dixon and Grace, 1984; Gutiérrez et al., 1994), or found transpiration to be insensitive to U (Komatsu et al., 2006).

Here we analyze the effect of U on transpiration through measurements and by quantifying the response to U of estimated resistances in the water vapor pathway, focusing on the reduced expression:

$$R_T = R_A + R_{bl} + R_{stom} \quad (4)$$

Many of the studies on transpiration response to U were performed in wind tunnels (Caldwell, 1970; Chu et al., 2009; Dixon and Grace, 1984; Grace and Russel, 1982; Kitaya et al., 2004). However, results of wind tunnel experiments may differ from those of field experiments because they do not account for the effect of neighboring crowns (shelter effect) on the coupling between stomata

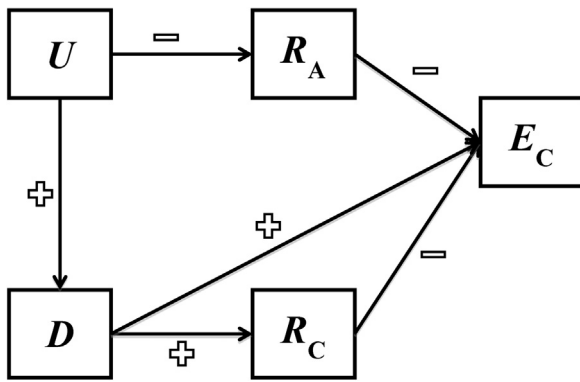


Fig. 1. Diagram showing the effect of wind velocity, U , on transpiration, E_C in a big-leaf framework. The sign represents the relationship between the two parameters at the arrow tip and tail (D – vapor pressure deficit; R_A – aerodynamic resistance; R_C – canopy resistance). For example, an increase in U results in a decrease in R_A because of the negative relationship between U and R_A and, in turn, the decrease in R_A leads to an increase in E_C because of the negative relationship between R_A and E_C .

and atmosphere in forest stands. In situ experiments were performed to overcome the limits of wind tunnels, but most have focused on crops and other small statured canopies; eight of the ten in situ studies we found were made on such canopies (Table 1). The response of these canopies to U may differ from that of forests canopies which tend to be rougher, and thus better coupled to the atmosphere (Jones, 1992). Furthermore, the aerodynamic roughness of forest canopies depends on tree height, stand density, tree size distribution, and LAI, all of which change with stand age (Chapin, 2002; Perry, 2008). Thus, if R_A is actually large, it would require quantification for each combination of stand characteristics before estimating biosphere–atmosphere exchanges.

As LAI increases, R_C decreases, improving the canopy–atmosphere coupling (Granier and Bréda, 1996; Granier et al., 2000; Katsoulas et al., 2007), yet the increase in canopy leaf area may impede air movement, increasing R_A while decreasing D within the canopy. At a given canopy LAI, as U increases, the accompanying decrease of R_A allows for frequent replacement of moist canopy air with dry air from aloft, potentially driving more transpiration, yet the drier air may force stomatal closure and increase R_{stom} (Fig. 1). Thus, the effect of U on tree and canopy transpiration represents a balance among interacting processes and may not be simple to predict. Transpiration could therefore decrease with increasing U if R_C increases as a result of stomatal response to D , overwhelming both the decrease of R_A as well as the increase of D , the driving force for transpiration.

The goals of this study were to improve understanding of the effect of U on transpiration. We assessed the response to U of

sap flux density (J_S), a proxy of transpiration, of six co-occurring species, ranging in shade tolerance, leaf dimensions and leaf clumping, in a mature broadleaved deciduous forest, testing the following hypotheses: **(H1)** transpiration will increase with U as a result of the combination of direct and indirect effects on resistances and forces above and within the canopy; **(H2)** R_A has a small effect on R_T meaning that R_T is dominated by the components of R_{sur} (mostly R_C , in-turn dominated by R_{stom}); thus **(H3)** the response of E_C to U will be determined by the effects of increases in D with U on R_{stom} and the increase in the driving force, rather than by the direct effect of U on R_A and R_{bl} .

2. Materials and methods

2.1. Site description and data collection

The study was conducted at the Duke Forest hardwood site, North Carolina (36°58' N, 79°05' W) from 2002 to 2005. The maximum tree age of the hardwood stand is ~100 years. Mean and maximum canopy heights of the stand are 25 m and 35 m, respectively (Pataki and Oren, 2003). A total of 40 trees of the six dominant species, yellow (tulip) poplar (*Liriodendron tulipifera*), sweetgum (*Liquidambar styraciflua*), hickory (*Carya tomentosa*), white oak (*Quercus alba*), swamp chestnut oak (*Q. michauxii*), and willow oak (*Q. phellos*), were selected for sap flux measurement (Table 2). J_S was measured with 20 mm long Granier-type heat dissipation sensors (Granier, 1987). The Granier-type technique has a potential to produce bias of J_S due to its sensitivity to physical factors including thermal conductivity of sapwood, temperature gradients within sapwood, and variation in J_S with sapwood depth (Tatarinov et al., 2005; Wullschlegel et al., 2011). However, using the same data, Oishi et al. (2010) showed that J_S produced similar estimates of latent heat flux when properly scaled to the site's eddy-covariance footprint. The measurement depths varied from 0–20 to 0–60 mm based on the expected depth of sapwood; sensors generating values at least two standard deviations apart from values collected by others in the same species and measurement depth were excluded to avoid potential underestimation caused by excessive penetration into heartwood (Lu et al., 2004).

T_a and relative humidity were measured at two-thirds of the mean canopy height with a temperature/RH probe (HMP35C, Campbell Scientific, Logan, UT, USA), and used to calculate D (Campbell and Norman, 1998). Photosynthetically active radiation (PAR) was measured above the canopy using a quantum sensor (LI-190SA, LI-COR, Lincoln, NE, USA). Volumetric soil moisture (θ) was measured in 12 locations, at 5–10 cm and 20–25 cm in each location, using ThetaProbe sensors (ML2x, Delta-T Devices, Cambridge, UK). U was measured with a tri-axial sonic anemometer installed above the canopy at 39.8 m (CSAT3, Campbell Scientific, Logan, UT,

Table 2
Characteristics of sampled trees at the Duke Forest hardwood site. Standard deviations are shown in parentheses.

Species	DBH ^a	LAD ^b	J_S ^c	A_S ^d	SLA ^e	d^f	Number of trees
<i>L. tulipifera</i>	39.3 (14.96)	0.017	42.36(15.21)	2.37	161.8	8.3	10
<i>L. styraciflua</i>	36.58(12.16)	0.020	29.44(13.10)	2.09	102.4	8.1	10
Other diffuse porous	–	–	–	1.48	173.3	–	–
<i>C. tomentosa</i>	34.12(20.86)	0.062	36.32(15.30)	2.11	131.9	4.7	5
<i>Q. alba</i>	32.34(18.46)	0.018	42.30(12.61)	0.52	112.2	5.4	5
<i>Q. michauxii</i>	33.68(16.80)	0.014	37.55(10.40)	0.59	128.9	6.8	5
<i>Q. phellos</i>	49.48 (9.02)	0.003	43.88(10.28)	0.27	101.9	1.1	5
Other ring porous	–	–	–	0.60	–	–	–

^a Diameter at breast height (cm).

^b Leaf area density ($m^2 m^{-3}$).

^c Average sap flux density ($g H_2O m^{-2} s^{-1}$) from selected data.

^d Sapwood area ($cm^2 m^{-2}$) weighed based on radial profiles of J_S (Oishi et al., 2008).

^e Specific leaf area ($cm^2 g^{-1}$).

^f Leaf characteristic dimension (cm; Harlow et al., 1996).

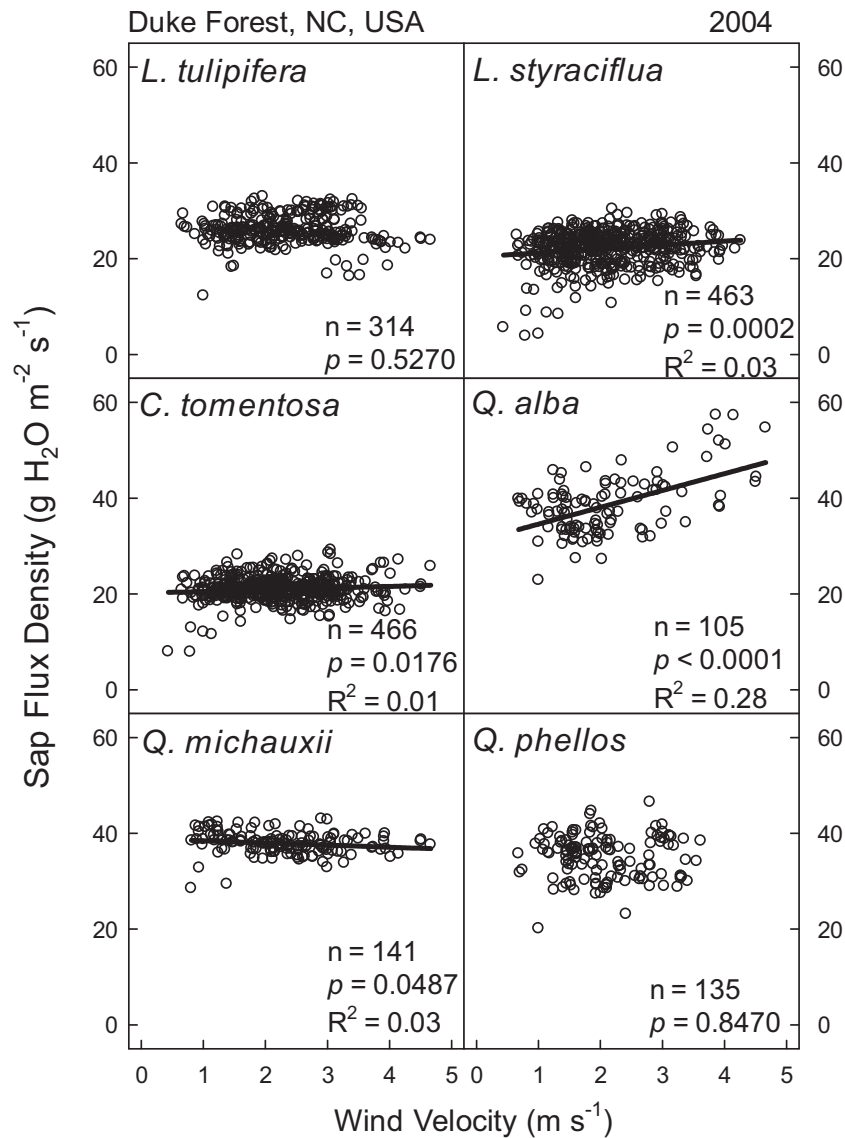


Fig. 2. The effect of wind velocity on sap flux density of six species in the Duke Forest hardwood site. Solid lines represent significant relationships at $p < 0.05$.

USA). More details on the site and measurements are given in previous publications on the stand (Oishi et al., 2008; Pataki and Oren, 2003; Stoy et al., 2006). Thirty-minute averages of all the measurements were stored on data loggers (CR23X, Campbell Scientific, Logan, UT, USA).

2.2. Data preparation

To minimize the effects of meteorological variables other than U , we selected data during times in which leaf area index is reasonably stable (DOY 120–275, excluding the short period of fast leaf expansion and the long period of leaf abscission and shedding) and conditions where transpiration is not likely limited by T_a ($\geq 25^\circ\text{C}$) and θ ($\geq 0.2\text{ m}^3\text{ m}^{-3}$). Data under these conditions were further partitioned to four ranges of light (PAR, at $500\text{ }\mu\text{mol m}^{-2}\text{ s}^{-1}$ up to $1500\text{ }\mu\text{mol m}^{-2}\text{ s}^{-1}$, and one class for data under $\text{PAR} \geq 1500\text{ }\mu\text{mol m}^{-2}\text{ s}^{-1}$, on which we concentrated most of the analyses for reasons that will become clear later).

J_s varies greatly among sensors (Granier, 1987). Typically, before estimating species or stand scale transpiration, missing data is gap-filled to avoid the effect on the average J_s of single sensors

dropping in and out of the data set. However, we chose a conservative approach, relying on only operating sensors, selecting the largest number of sensors for each species resulting in at least 3000 half-hourly, species level averages over the study period, all produced by the same three to five sensors. The 2003 J_s data of *Q. michauxii* were excluded from analysis because the values were unstable. When calculating stand-scale E_c , missing average J_s of any species ($\sim 54\%$ of potential values) resulted in missing half-hourly stand-scale E_c . To gap-fill the missing average J_s , a power function ($y = a + bx^n$) was fitted among species when average J_s of at least three species were available ($\sim 64\%$ of the time under the selected conditions). If the data of more than three species at a given half-hour were missing, the data were neither gap-filled nor used for E_c calculation; otherwise, missing data of a species were gap-filled using data of another species with the best fit in terms of R^2 value. Approximately 50% of the missing data were gap-filled. The performance of the gap-filling procedure was evaluated by comparing the data unused in developing the gap-filling functions ($\sim 36\%$ of the time, when data for only one or two species were available) and estimates based on the function developed for each species. The estimated values were well fitted with observed ones ($R^2 > 0.89$; slope ranged from 0.88 to 1.10).

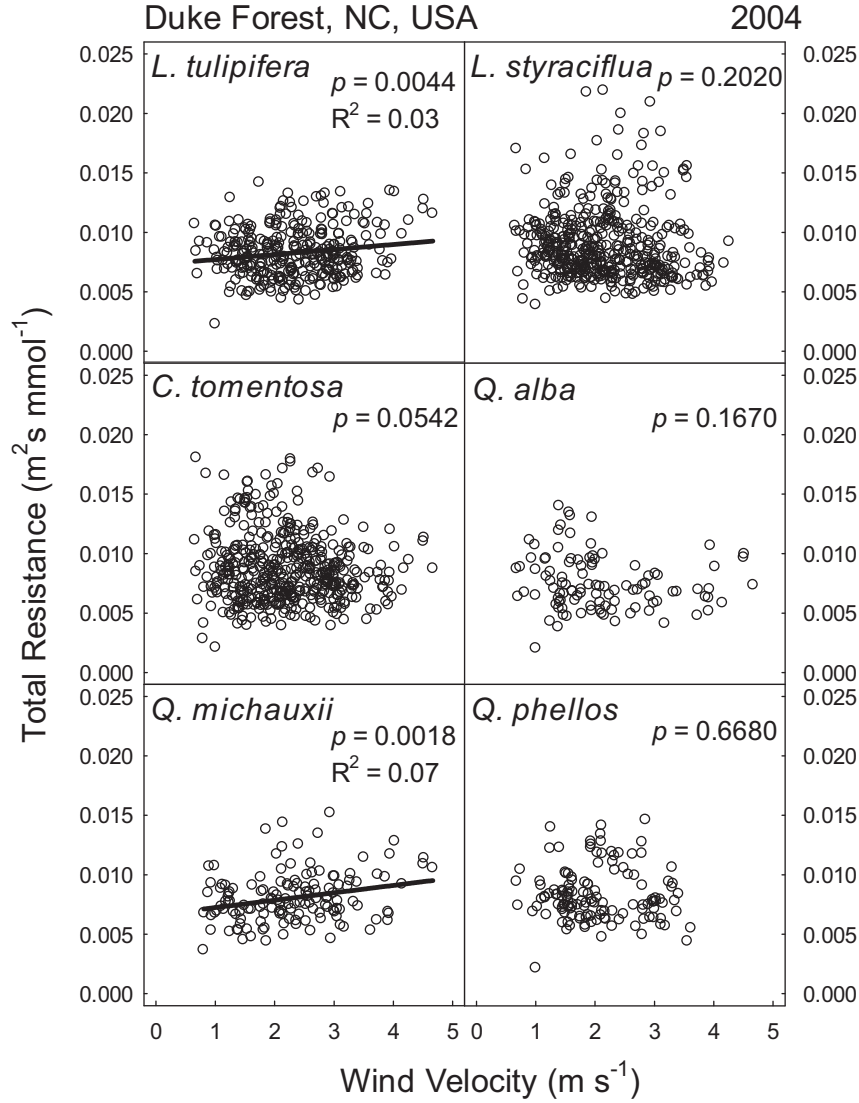


Fig. 3. The relationship between wind velocity and total resistance. Solid lines represent significant relationships at $p < 0.05$.

R_A in the air layer above the canopy surface was estimated by Monin–Obukhov similarity theory (Campbell and Norman, 1998):

$$R_A = \frac{[\ln((z-d)/z_{om}) + \Psi_M][\ln((z-d)/z_{oh}) + \Psi_H]}{k^2 \hat{\rho} U} \quad (5)$$

$\times (\text{m}^2 \text{ s mmol}^{-1})$

$$\Psi_H = \begin{cases} 6 \ln(1 + \zeta), & \text{for stable flow} \\ -2 \ln \left[\frac{1 + (1 - 16\zeta)^{1/2}}{2} \right], & \text{for unstable flow} \end{cases} \quad (6)$$

$$\Psi_M = \begin{cases} \Psi_H, & \text{for stable flow} \\ 0.6\Psi_H, & \text{for unstable flow} \end{cases} \quad (7)$$

where $\hat{\rho}$ is the molar density of air (mmol m^{-3}), k ($=0.4$) is the von Karman constant, U is the wind velocity (m s^{-1}), z is the height of measured wind velocity (m), d is the zero plane displacement (m), z_{om} is the momentum roughness length (m), z_{oh} is the scalar roughness length for heat transfer (m), Ψ_M and Ψ_H are the profile diabatic

correction factors for buoyant instability, ζ is the atmospheric stability. The parameters d , z_{om} , and z_{oh} were taken as fixed values, 0.67, 0.1, and 0.02 of the canopy height in this study, making the effect of the surface roughness length constant during the growing season. A 10% variation of each of the three parameters changed the maximum of calculated values of R_A by 2.8, 1.8, and 1.7%, supporting the assumption of constant parameters.

R_C was obtained by subtracting R_A from the total resistance (R_T), which was calculated from a rearranged Penman–Monteith equation after Blanken and Black (2004):

$$R_T = \frac{s}{\gamma} \beta R_A + \frac{\rho c_p D}{\hat{\rho} \gamma \lambda E_C} \quad (\text{m}^2 \text{ s mmol}^{-1}) \quad (8)$$

where s is the slope of the saturation vapor pressure with T_a (kPa K^{-1}), γ is the psychrometric constant (kPa K^{-1}), β is the Bowen ratio (ratio of sensible to latent heat fluxes), ρ is the density of air (kg m^{-3}), c_p is the specific heat of air ($\text{J kg}^{-1} \text{ K}^{-1}$), D is the vapor pressure deficit (kPa), λ is the latent heat of vaporization (J g^{-1}), and E_C is the canopy transpiration ($\text{g H}_2\text{O m}^{-2} \text{ (ground area) s}^{-1}$). Typical values of β for temperate forests (lower in deciduous than in conifer forests) ranged from 0.4 to 0.8 (Novel, 2005; Wilson et al., 2002). Since a change of β by 0.1 within the range resulted in only 2.5% difference in R_T under the selected conditions in this study, β was

assumed to be constant (0.5). E_C was estimated with J_S from outer sensors (depth of 0–20 mm; $\text{g H}_2\text{O m}^{-2}$ (sapwood area) s^{-1}) and cross-sectional sapwood area per unit ground area (A_S ; $\text{cm}^2 \text{m}^{-2}$) weighed using radial profiles of J_S . The weighed A_S of each species was obtained from allometric relationship between diameter at breast height and sapwood area (Oishi et al., 2008, 2010).

To assess the importance of R_{bl} in this stand, species mean R_{bl} values were calculated using an equation for the resistance of vapor transport in forced convection (Campbell and Norman, 1998), which dominates over free convection in this canopy (Stoy et al., 2006):

$$R_{bl} = \frac{\nu^{1/6}}{0.664 \rho D_v^{2/3}} \sqrt{\frac{d_l}{u}} = 0.0068 \sqrt{\frac{d_l}{u}} (\text{m}^2 \text{s}^{1/2} \text{mmol}^{-1}) \quad (9)$$

where ν is the kinematic viscosity of air ($\text{m}^2 \text{s}^{-1}$), D_v is the molecular diffusivity of water vapor in air ($\text{m}^2 \text{s}^{-1}$), d_l is the leaf characteristic dimensions (0.011–0.083 m; Table 2), and u is the vertically averaged U within canopy (m s^{-1}). A constant value of $0.0068 \text{ m}^2 \text{s}^{1/2} \text{mmol}^{-1}$ was used for $\nu^{1/6} (0.664 \rho D_v^{2/3})^{-1}$ (Campbell and Norman, 1998). u was estimated based on U profiles within canopy and LAI profile when canopy leaf area was at its stable seasonal plateau (DOY 154):

$$u = \frac{\sum_{h=1}^H U(h) \text{LAI}(h)}{\sum_{h=1}^H \text{LAI}(h)} \quad (10)$$

where $U(h)$ and $\text{LAI}(h)$ are U and LAI at a height h within canopy, and H is the canopy height (m). The U profiles were calculated as (Campbell and Norman, 1998):

$$U(h) = u_0 e^{a((h/H)-1)} \quad (11)$$

where u_0 is the U at the top of the canopy and a is an attenuation coefficient (assumed to be 2.5). The LAI profile was estimated from measurements at ~ 2 m intervals from 1 to 45 m along a tower using an optical plant canopy analyzer (LAI-2000, LI-COR, Lincoln, NE, USA; Stoy et al., 2005).

Also, to understand the coupling between the air within and above the canopy volume, the decoupling coefficient for the canopy (Ω) was considered. In this study Ω was calculated using Jarvis and McNaughton (1986):

$$\Omega = \frac{(s/\gamma) + 1}{(s/\gamma) + 1 + (R_C/R_A)} \quad (12)$$

Ω approaches zero when the ratio R_C/R_A approaches infinity, a condition of perfect coupling between the air in the canopy volume and the bulk air above the canopy (i.e. very efficient mass and energy exchange), resulting in essentially identical T_a and humidity in these two locations. If R_A is much larger than R_C , Ω approaches one, reflecting the isolation of the air in the canopy volume from that aloft.

3. Results and discussion

3.1. Sap flux density

We first examined the relationship between U and J_S . We used a subset of the available data to minimize the effects of other variables including light, temperature, and soil water. The means (and ranges) of T_a , PAR, θ , D , and U of the selected data were 29.0 (25.0–35.1) °C, 1662 (1500–2047) $\mu\text{mol m}^{-2} \text{s}^{-1}$, 0.30 (0.20–0.50) $\text{m}^3 \text{m}^{-3}$, 1.69 (0.09–3.09) kPa, and 2.16 (0.44–5.09) m s^{-1} , respectively. Although the size of the subset is only 5.3% of total daytime of the stable LAI period (but see below), the range of U includes 98.3% of the range of U during daytime hours of the growing season above this stand. The average

J_S of *L. tulipifera*, *L. styraciflua*, *C. tomentosa*, *Q. alba*, *Q. michauxii*, and *Q. phellos* after the filtering were 26.0, 22.1, 20.9, 38.5, 37.8, and 35.0 $\text{g H}_2\text{O m}^{-2}$ (sapwood area) s^{-1} , respectively.

Fig. 2 shows the relationship between 30 min average J_S of the six species and U during the growing season. Over the wide range of U , J_S of *L. tulipifera* and *Q. phellos* did not change significantly. For *L. styraciflua*, *C. tomentosa*, *Q. alba*, and *Q. michauxii*, the relationships were weak ($R^2 = 0.01$ – 0.28) and inconsistent in sign. The sensitivity of J_S to U can be influenced by different leaf size among species because specific leaf dimension (i.e. leaf width in the direction of the wind) changes the leaf boundary layer thickness (Campbell and Norman, 1998). However, no effect of leaf size on the J_S – U sensitivity was observed among species ($p = 0.796$).

Observations of the relationship between U and sap flux (or tree transpiration) were diverse. For example, increasing U by up to 2 m s^{-1} by partially harvesting a boreal stand composed of *Picea glauca*, *P. balsamifera*, and *Butula papyrifera* increased transpiration per unit leaf area (Bladon et al., 2006), but sap flux of hybrid poplar (*Populus nigra* L. $\times$ *P. maximowiczii* A. Henry) was negatively correlated to U (Zalesny et al., 2006). Komatsu et al. (2006) found no difference of sap flux of *Cryptomeria japonica* under a given D at all $U > 0.7 \text{ m s}^{-1}$. In a study of 7-year-old *Eucalyptus globulus* the ratio of sap flux of edge trees to that of trees far from the edge increased with U when R_C of edge trees was sufficiently low to cause leaf temperature to equal T_a (Taylor et al., 2001). Similarly at the individual crown scale, simulated transpiration of 20-year-old *Juglans regia* foliage increased slightly as U increased from 1 to 5 m s^{-1} in shaded leaves but was insensitive to U in sunlit leaves that had higher leaf temperature than T_a (Daudet et al., 1999). The results of these studies demonstrate the complex effect U can have on stand transpiration.

3.2. Total, aerodynamic, and canopy resistance

In order to examine how U affected resistances, we first calculated R_T for each species from its J_S and A_S , as if the stand was entirely composed by that species (Fig. 3). The mean R_T values of each species (*L. tulipifera*, *L. styraciflua*, *C. tomentosa*, *Q. alba*, *Q. michauxii*, and *Q. phellos*) were 0.83×10^{-2} , 0.89×10^{-2} , 0.87×10^{-2} , 0.75×10^{-2} , 0.81×10^{-2} , and $0.83 \times 10^{-2} \text{ m}^2$ (ground area) s mmol^{-1} , respectively. R_T weakly increased with U for two species in the stand (*L. tulipifera* and *Q. michauxii*; Fig. 3).

Stand-level R_T based on the actual species composition was calculated using sap flux-scaled canopy transpiration (E_C). E_C was unrelated to U in three PAR intervals (0–500, 500–1000, and 1000–1500 $\mu\text{mol m}^{-2} \text{s}^{-1}$) when analyzed along the entire range of U found in each interval ($p = 0.0746$, 0.3710 , and 0.0707 ; Fig. 4a). In the highest PAR interval ($\geq 1500 \mu\text{mol m}^{-2} \text{s}^{-1}$), a weak positive relationship was observed between E_C and U ($p = 0.0069$, $R^2 = 0.02$). However, the increase of E_C associated with ~ 7 -fold increase of U was only 6.5%. Thus, we must reject **H1**: under most conditions, stand transpiration did not increase with U .

Although R_T increased with U for two species, the relationship between stand-level R_T and U was not significant ($p = 0.3650$; Fig. 4b). Because of the observed lack of or limited sensitivity of E_C to U in any PAR range, we concentrate our analyses on only the one set of conditions in which no environmental variable was limiting, thus reducing the variation in the calculated resistances. The direct effect of U on water transport through the air layer above the canopy is reflected in R_A . The average R_A of this stand was $4.62 \times 10^{-4} \text{ m}^2$ (ground area) s mmol^{-1} , within the range of values from other broadleaved forest stands (Köstner et al., 1992; Kumagai et al., 2004; Lindroth, 1993; Pereira et al., 2010). The estimated values of R_A over the canopy decreased from 1.31×10^{-3} to $1.95 \times 10^{-4} \text{ m}^2$ (ground area) s mmol^{-1} as U increased from ~ 0.7 to $\sim 4.7 \text{ m s}^{-1}$ (Fig. 4c). However, consistent with **H2**, R_A was a

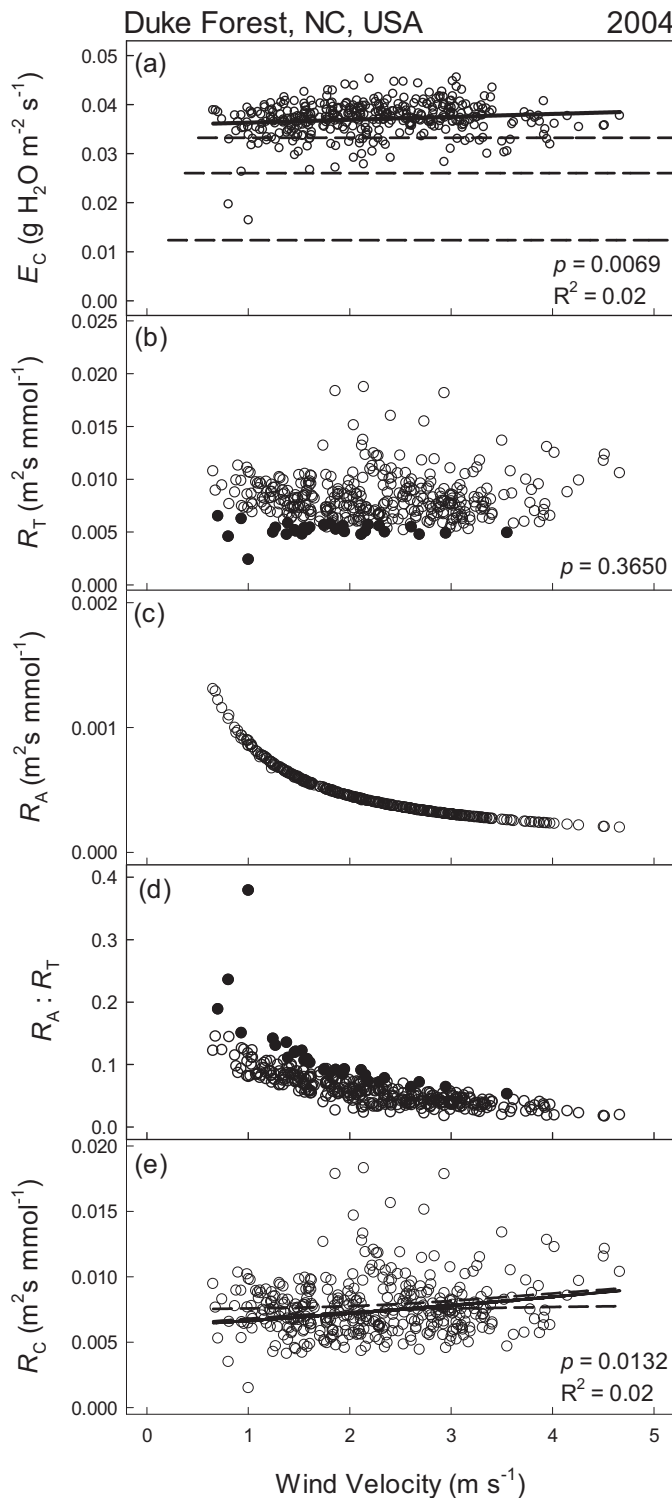


Fig. 4. Changes in stand-level transpiration and resistances with increasing wind velocity (U). (a) Canopy transpiration (E_c); dashed lines from the bottom up represent light conditions of $500 \mu\text{mol m}^{-2} \text{s}^{-1}$ intervals from 0 to $1500 \mu\text{mol m}^{-2} \text{s}^{-1}$, while the data points represent conditions of $>1500 \mu\text{mol m}^{-2} \text{s}^{-1}$. Only one PAR interval ($>1500 \mu\text{mol m}^{-2} \text{s}^{-1}$) showed significant yet very weak correlation (solid line). (b) total resistance (R_T); filled circles represent R_T when vapor pressure deficit (D) is low ($<1 \text{ kPa}$), (c) aerodynamic resistance (R_A), (d) proportion of R_A to R_T ; filled circles represent the proportion when D is low ($<1 \text{ kPa}$), and (e) canopy resistance (R_C) with 95% confidence intervals for individual estimates (dashed lines) and expected R_{stom} based on the relationship between R_{stom} and vapor pressure deficit (D), and between D and U (solid line; see text).

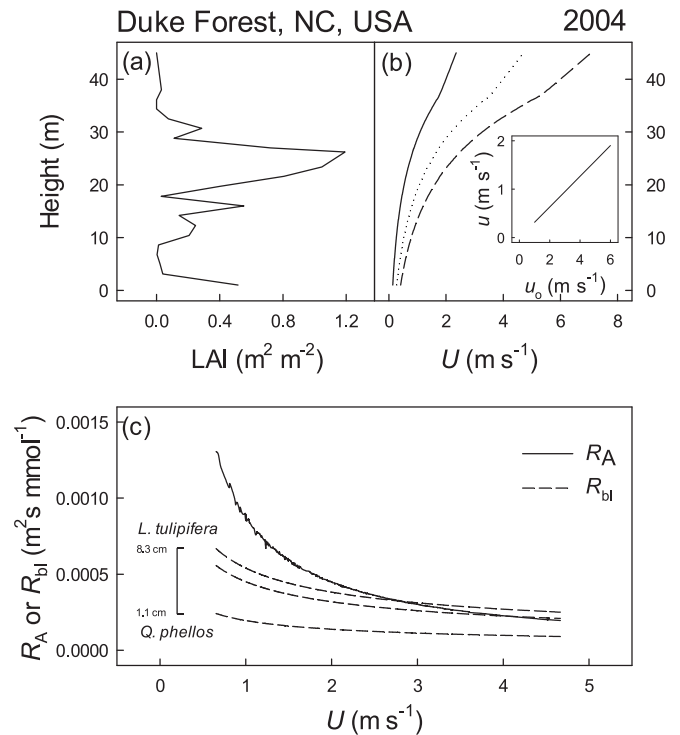


Fig. 5. (a) Leaf area index (LAI) profile (DOY 154). (b) Calculated wind velocity (U) profiles above and within the canopy when U at measurement height (39.8 m) are 2 (solid line), 4 (dotted line), and 6 m s^{-1} (dashed line). Inset figure shows the relationship between U above canopy (u_0) and average U within canopy weighted by leaf area profile (u). (c) Aerodynamic resistance (R_A ; solid line) and leaf boundary layer resistance (R_{bl} ; dashed lines) estimated using U and leaf characteristic dimensions of the six species, ranged from 1.1 cm for *Q. phellos* to 8.6 cm for *L. tulipifera* (see Table 2); the dashed lines from the top represent R_{bl} estimates of *L. tulipifera*, weighed average, and *Q. phellos*.

relatively small component of R_T , decreasing from 15% to 1.6% with increasing U except three half-hourly values with 18.8, 23.5 and 37.8% due to exceptionally low D (0.10, 0.31 and 0.97 kPa; Fig. 4d). When D was low, R_T was generally lower under all U condition, and thus the ratio of R_A to R_T increased somewhat with decreasing D (see Fig. 4b and d).

R_C , calculated by subtracting R_A from R_T , increased with U in the stand ($p=0.0132$; $R^2=0.2$; Fig. 4e). The mean R_C of the stand was $7.61 \times 10^{-3} \text{ m}^2 (\text{ground area}) \text{s mmol}^{-1}$, similar to values observed in other broadleaved forest stands (Blanken and Black, 2004; Hinckley et al., 1994). In a *Coffea arabica* plantation, R_C normalized by PAR also increased with increasing U from 1 to 2.5 m s^{-1} (Gutiérrez et al., 1994), as it did, albeit weakly, in a *Zea mays* field (Irmak and Mutiibwa, 2010).

We assessed the contribution of R_{bl} to R_C modeling the wind velocity profile down the canopy based on leaf-area profile and characteristic leaf dimensions (Fig. 5a and b). The maximum values of R_{bl} of the six species ranged from 2.76×10^{-4} to $7.64 \times 10^{-4} \text{ m}^2 (\text{ground area}) \text{s mmol}^{-1}$, an order of magnitude smaller than R_C over much of the range of U in this stand (Figs. 4e and 5c). Indeed, R_{bl} was less than half of R_A at low U , becoming similar to R_A only at very high U when the sum of both resistances was a very small component of the total. Thus, R_{stom} is much greater than R_{bl} and can be assumed equivalent to R_C . We test this assumption when we address H3 in the last section.

The small absolute decrease of R_A with increasing U , and the low ratio of R_A to R_T at our stand indicate that the stand was generally well-coupled to the atmosphere (mean $\Omega=0.22$; Fig. 6a) despite its high tree density (~ 866 total trees per hectare), extensive vertical distribution of foliage (~ 685 of ~ 866 trees were in the

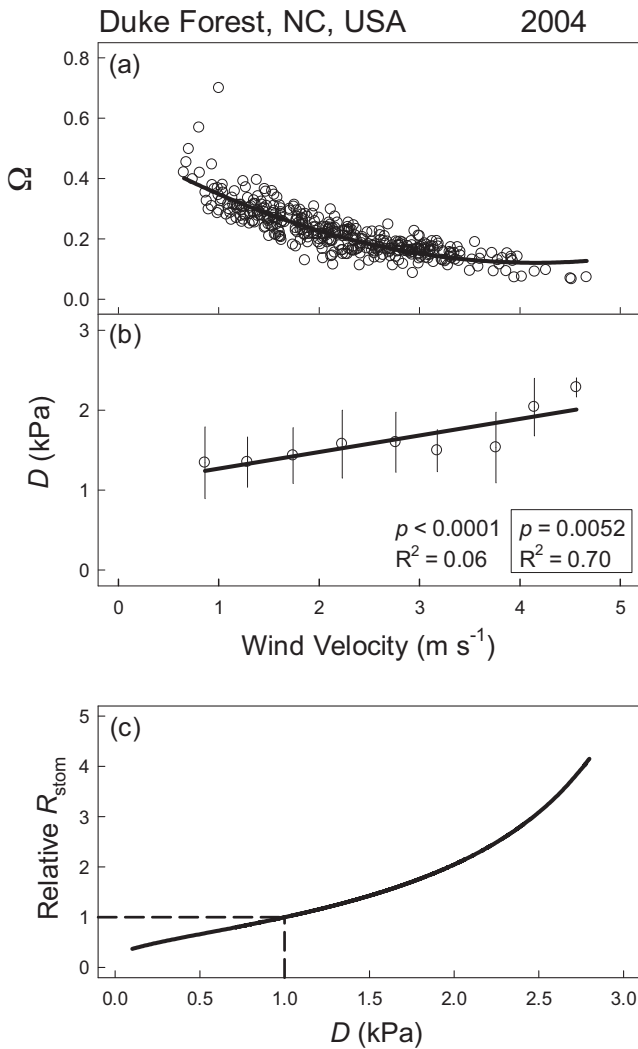


Fig. 6. (a) Responses of the decoupling coefficient (Ω) to wind velocity (U). (b) Responses of vapor pressure deficit (D) to U (data are mean vapor pressure deficit grouped in U intervals of 0.5 m s^{-1} ; p -value and R^2 within the box were calculated from mean values rather than individual data). (c) The theoretical, hydraulically controlled increase of stomatal resistance (R_{stom}) with D , normalized by the resistance at $D = 1 \text{ kPa}$.

sub-canopy) and high LAI ($\sim 7 \text{ m}^2 \text{ m}^{-2}$). Transpiration in a stand with low Ω value is largely influenced by air humidity and stomatal response rather than available energy (Jarvis and McNaughton, 1986; Lindroth, 1993), which in our analysis was set to a limited range (approximately $1500 \leq \text{PAR} \leq 2050 \mu\text{mol m}^{-2} \text{ s}^{-1}$, and $25 \leq T_a \leq 35^\circ\text{C}$). Thus, we assessed further the concurrent variation of D and R_{stom} .

3.3. Vapor pressure deficit and its effect on resistances

D functions as a driving gradient for water vapor exchange through stomata, where carbon assimilation driven by CO_2 concentration gradient occurs concurrently. Understanding the effect of U on transpiration is complicated because wind affects not only R_A and R_{bl} , but also D , and thus R_{stom} (Fig. 1; Table 1; Dixon and Grace, 1984). Gutiérrez et al. (1994) suggested that increased leaf-to-air D at the leaf surface with U generated a positive relationship between normalized R_C and U , but evidence of such relationships is rare. In this study, the slight increase in canopy-atmosphere coupling of the stand was sufficient to produce a significant positive relationship ($p < 0.0001$) between U and D (Fig. 6b). Grace et al. (1975)

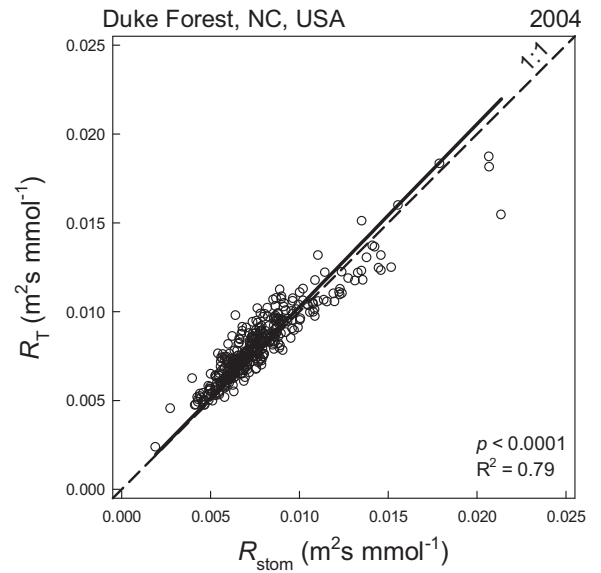


Fig. 7. The relationship between predicted stomatal resistance (R_{stom}) and total resistance (R_T) estimated from sap flux density and associated measurements. R_{stom} was calculated using equation 10, with $R_{\text{Sref}} = 0.51 \times 10^{-2} \text{ m}^2 (\text{ground area}) \text{ s mmol}^{-1}$ and $m = 0.74$. Solid line and dashed line represent a linear fit without intercept and 1:1 line.

estimated that as U increased from 0.1 to 1.0 m s^{-1} , the effective D at leaf surface increased by 38%. Within the range of our data, the increase of D within the canopy volume averaged $\sim 11\%$ per 1 m s^{-1} increase in U , with greater increases at the lower end. At low D , its increase tends to drive higher canopy transpiration (Daudet et al., 1999; Kupper et al., 2011). However, further increases in D lead to stomatal closure (decrease in leaf stomatal conductance) such that transpiration does not increase through the entire range of D commonly observed in temperate forests (Oishi et al., 2010; Oren and Pataki, 2001; Oren et al., 1999; Pataki and Oren, 2003). Phillips and Oren (2001) explained the small variation observed in canopy transpiration of *Pinus taeda* during periods of daytime $D > 0.9 \text{ kPa}$ based on compensation of the direct (positive) and indirect (negative) effects of D . Indeed, Roberts (1983) suggested the negative feedback between R_{stom} and D as the reason transpiration is a conservative hydrological process in forests.

More common are studies showing increased R_{stom} or R_C of broadleaved stands with increased D (Blanken and Black, 2004; Clausnitzer et al., 2011; Granier et al., 2000; Irmak and Mutiibwa, 2010; Jonard et al., 2011; Köstner et al., 1992; Oren and Pataki, 2001; Pataki and Oren, 2003; Schäfer et al., 2000). The response of R_{stom} to D can be estimated as (Lohammar et al., 1980; Oren et al., 1999):

$$R_{\text{stom}} = \frac{R_{\text{Sref}}}{1 - m \ln(D)} \quad (13)$$

where R_{Sref} is a reference resistance at $D = 1 \text{ kPa}$, m is the slope of the relationship between reference conductance ($1/R_{\text{Sref}}$) and the sensitivity of stomatal conductance to D ($-dG_{\text{stom}}/d \ln(D)$). This response is consistent with hydraulic theory and reflects the decreased sensitivity to D of R_{stom} with increasing R_{stom} at low D (Katul et al., 2009). We estimated R_{Sref} and m from coefficients of linear regression between canopy conductance (a reciprocal of R_C) and the natural log of D . The estimated value of m , 0.74, was higher by about 9% than that estimated in similar forest stands in the area (Oren and Pataki, 2001). Based on R_{Sref} and m we estimated the expected increase of R_{stom} with D (Fig. 6c). Based on this analysis, as D increases from 0.1 to 3.0 kPa, R_{stom} should increase about ten-fold, increasing two-fold between 1.0 and 2.0 kPa, similar to observation in a 35-year-old oak stand (Granier and Bréda, 1996).

We examined the expected response of R_{stom} to U by converting D in equation 13 to U based on the relationship between D and U (from Fig. 6b). Estimated R_{stom} increased 38% over the range of U and fell close to or within the 95% confidence intervals of the relationship between R_C and U (Fig. 4e). As a result of the small contribution of R_A to R_T , the variation of R_T was also highly correlated ($p < 0.0001$, $R^2 = 0.79$) with the response of R_{stom} to D ; the slope was very close to unity, with R_T averaging only 2.8% higher than the expected values of R_{stom} (Fig. 7). Thus, we failed to reject **H3**, concluding that the direct effect of D , acting as a driving gradient for plant transpiration, relative to its indirect effect, inducing stomatal closure (Fig. 1), dominates the response of transpiration to U . The average D , even in calm wind conditions of less than 1 m s^{-1} , was $>1 \text{ kPa}$ in this study, suggesting that much of the time the direct effect of D was offset by R_{stom} response to D , resulting in a constant E_C over the observed range of U .

4. Conclusion

This study assessed the responses of J_S (as a surrogate for transpiration) to U in a mixed hardwood stand under optimal meteorological condition for transpiration based on T_a and θ . J_S of the six species were insensitive or weakly responded to increasing U ; stand-level transpiration appeared not to significantly change with U (**H1**). The canopy of the broadleaved stand was well-coupled to the atmosphere, and thus had a low R_A . Further decrease in R_A with increasing U was readily offset by an increasing R_C (**H2**). The increase in R_C was consistent with the theoretical increase in R_{stom} with D – which in turn increased with U – keeping transpiration constant despite the increase of the driving force (**H3**). These results suggest that in similar forests the effect of U on canopy transpiration is negligible. However, due to its effect on D within the canopy and thus R_{stom} , U could have an important influence on CO_2 uptake.

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