



Functionally relevant climate variables for arid lands: a climatic water deficit approach for modelling desert shrub distributions

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ABSTRACT

Aim We have three goals. (1) To develop a suite of functionally relevant climate variables for modelling vegetation distribution on arid and semi-arid landscapes of the Great Basin, USA. (2) To compare the predictive power of vegetation distribution models based on mechanistically proximate factors (water deficit variables) and factors that are more mechanistically removed from a plant's use of water (precipitation). (3) To quantify the climate gradients that control shrub distributions in a cold desert environment.

Location The central basin and range ecoregion of the western USA (36–42°N).

Methods We used a modified Thornthwaite method to derive monthly water balance variables and to depict them using a water balance climograph. Eighteen variables were calculated from the climograph, representing different components of the seasonal water balance. These were used in boosted regression tree models to derive distribution models for 18 desert shrub species. The water balance approach was compared with an approach that used bioclimatic variables derived from the PRISM (Parameter-elevation Relationships on Independent Slopes Model) climate data.

Results The water balance and bioclimatic approaches yielded models with similar performance in predicting the geographical distribution of most shrub species. Cumulative annual climatic water deficit was consistently the most important water balance variable for predicting shrub type distributions, although predictions were improved by the inclusion of variables that describe the seasonal distribution of water balance such as water supply in the spring, fall actual evapotranspiration, monsoonality and summer decline in actual evapotranspiration.

Main conclusions The water balance and bioclimatic approaches to species distribution modelling both yielded similar prediction accuracies. However, the water balance approach offers advantages over the bioclimatic approach because it is mechanistically derived to approximate physical processes important for plant growth.

Keywords

Climatic water deficit, climograph, cold desert shrubs, gradient analysis, Great Basin, mechanistic model, PRISM, species distribution modelling, water balance, western USA.

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INTRODUCTION

A long tradition of vegetation modelling originates from gradient analysis studies. It is well established that vegetation

distribution responds to gradients of temperature, moisture, and to a lesser degree, soil nutrients and geology (e.g. Whittaker & Niering, 1965; Franklin, 1995). However, in arid and semi-arid regions of the world, water relations are expected

to play the dominant role in structuring plant communities. In mountainous arid regions, precipitation and temperature often vary consistently and inversely along the same gradient of elevation due to orographic and adiabatic effects. As a result, ecologists have characterized the vegetation of arid ecosystems as varying along a single moisture gradient (Whittaker, 1967) or have relied upon relationships with soils and topography to predict the distribution of plant species (West, 1983). The resulting models may be difficult to generalize to areas that have differing relationships between climate and topography. In climates in which the majority of precipitation falls outside the growing season, simple models may underestimate the role that soil serves to store water received during the wet season for use by plants during the growing season. Similarly, vegetation models relying upon topographic variables alone are not expected to be predictive under climate change scenarios, because topographic variables remain static while climatic variables change (Dyer, 2009). Because of the differing predictions for future precipitation and temperature regimes across space, a modelling approach that can functionally separate the water balance components (precipitation, evapotranspiration, soil water balance) may prove useful for predicting individual species responses to climate change.

Identifying species–environment relationships using the most direct causal variables is one of the fundamental challenges in ecology (Guisan & Zimmermann, 2000; Austin, 2002). The climatic water deficit (CWD) approach has been proposed as an ecologically functional description of climate influences that limit plant distributions (Stephenson, 1990, 1998). This approach uses water balance models and offers advantages for modelling species–environment relationships over traditional empirical ‘climate envelope’ models based purely on climate variables (precipitation and temperature) that use proxies for variables that are more functionally related to plant growth, including soil water content and soil water deficit. Directly including these critical variables in distribution models may increase their predictive capabilities and allow further insights regarding factors that limit establishment, growth and mortality (Piedallu *et al.*, 2013; Das *et al.*, 2013). At the level of the individual plant, what matters most is likely to be the amount of water in its rooting zone over the course of the time period when the plant most needs water, rather than the long-term average amount of precipitation at a regional level.

The temporal availability of moisture also is an important determinant of the distribution of major vegetation physiognomic types at local to continental scales (Stephenson, 1990). For example, both the Great Plains and Great Basin regions of the USA have similar annual precipitation and temperature regimes but the former is dominated by graminoids while the latter is dominated by shrubs. This difference is attributed to the lack of summer precipitation in the Great Basin which favours deep-rooted shrub species over graminoids (Paruelo & Lauenroth, 1996). The climatic water defi-

cit (CWD) approach, based upon a Thornthwaite water balance model (Lutz *et al.*, 2010), is well suited for representing the seasonal and temporal availability of moisture because it can be calculated at monthly time steps, tracking previous conditions and accounting for temporal lags between water availability and evaporative demand. The CWD approach can also include development of climographs (*sensu* Stephenson, 1998) that allow visual assessment of variables such as water supply and potential and actual evapotranspiration and provide a more nuanced view of vegetation–climate relationships. In contrast, the bioclimatic approach does not explicitly account for water storage in the soil or snowpack and considers temperature and precipitation as acting independently rather than interacting to affect water balance at a site.

Shrubs in the cold desert of the Great Basin of the western USA provide an excellent case for testing the water deficit approach and comparing it with traditional approaches using climate data based on precipitation and temperature. Desert shrubs have received less attention than trees with respect to vegetation–climate modelling (e.g. Lutz *et al.*, 2010; Crimmins *et al.*, 2011). Desert shrub distributions are strongly influenced by water availability and demand and vary with subtle changes in soil texture, soil depth and aspect (Franklin, 1995; Miller & Franklin, 2002, 2006), which can be represented in climate water deficit models. Here, we address three sets of research questions:

1. Do species distribution models developed from water deficit variables perform better in terms of model fit compared to distribution models developed from commonly used bioclimatic variables?
2. Are species distribution models developed from water deficit variables more transferable in space compared to distribution models developed from commonly used bioclimatic variables?
3. How are shrub species distributed along ecological gradients and which gradients are most predictive across species?

Previous exercises in species distribution modelling have demonstrated the predictive power of the annual water deficit and actual evapotranspiration for species distribution modelling (e.g. Stephenson, 1998; Lutz *et al.*, 2010; Piedallu *et al.*, 2013). We refine this general approach by developing new indices that more fully describe differences in demand and timing of water availability based on the water deficit climograph, and by incorporating these indices into our predictive models. We expect that variables quantifying annual levels of water availability and demand will be generally predictive, but that other variables that quantify the timing and seasonality of water availability and demand will help to distinguish among shrub species with differing phenology, seasonal patterns of water use and rooting depths. We compare models based on standard bioclimatic variables to models that use water deficit variables to determine whether a water deficit approach offers an advantage in predicting shrub type distributions. We expect the largest advantages of using

water deficit variables for lower-elevation shrub species that are most water-limited, and the least advantage for higher-elevation shrub species with greater water availability but that may be temperature-limited due to short growing seasons.

MATERIALS AND METHODS

Study area

Our study area is the entire Great Basin of the USA (451,112 km²) (Fig. 1), encompassing the majority of the Central Basin and Range province (Omernik, 1987). Land-form morphology is dominated by north–south trending fault block mountain ranges alternating with broad alluvial valleys. Elevations range from –83 to 4383 m. The climate is arid to semi-arid and average precipitation is 30.6 cm per year. Topographic effects on precipitation are strong, resulting in annual precipitation as high as 205 cm in mountainous areas and as low as 6.6 cm in the most arid sites. Winter months receive roughly three times as much precipitation as summer months, but, the eastern and southernmost portions of the study area are influenced by summer monsoonal storms.

Much of the region is dominated by deeply rooted shrub species. The lowest elevations are characterized by salt desert shrub communities and are dominated by shrub species such as shadscale (*Atriplex confertifolia* (Pursh) Nutt.), four-wing saltbush (*Atriplex canescens* (Pursh) Nutt.), winterfat (*Krascheninnikovia lanata* (Pursh) A. Meeuse & Smit), and black greasewood (*Sarcobatus vermiculatus* (Hook.) Torr). Sagebrush communities dominated by Wyoming big sagebrush (*Artemisia tridentata* Nutt. ssp. *wyomingensis* Beetle & Young), interspersed with perennial grasses, occur at middle elevations. At middle elevation sites with shallower or coarser soils, black sagebrush (*Artemisia nova* A. Nelson) dominates instead of Wyoming big sagebrush. At higher elevations mountain big sagebrush (*Artemisia tridentata* ssp. *vaseyana* (Rydb.) B. Boivin), low sagebrush (*Artemisia arbuscula* Nutt.) and antelope bitterbrush (*Purshia tridentata* Pursh) communities co-occur with single-needle pinyon pine (*Pinus monophylla* Torr. & Frém.) and Utah juniper (*Juniperus osteosperma* (Torr.) Little) (Tables 1 & 2).

Shrub species data

Presence–absence data were derived from Southwest Regional GAP Training Site Databases for Nevada and Utah (USDI, 2004a) for 18 of the most common desert shrub species (Table 1). The number of occurrence points per species ranged from 231 for four-wing saltbush to 3255 for Wyoming big sagebrush and species prevalence ranged from 0.6% to 8% (Table 2). Occurrence points falling within areas classified as riparian by the Southwest Regional GAP Analysis map (USDI, 2004b) were excluded, as were areas subject to recent fires based on the Monitoring Trends in Burn Severity

dataset (USDA, 2009) and the Nevada Fire History Database for 1981 to 2008 (USDI, 2008).

Environmental predictors

Climatic variables were derived from two primary sources. Eleven standard bioclimatic variables (Table 3) that are commonly used in bioclimatic envelope modelling (Hijmans *et al.*, 2005) were derived from gridded PRISM (Parameter–elevation Relationships on Independent Slopes Model) climate data (PRISM Group, 2007). Data from PRISM can substantially improve climate predictions in topographically heterogeneous regions compared to WorldClim and DayMet products (Daly *et al.*, 2008). Water deficit variables (Table 4) were derived from a water balance model (Lutz *et al.*, 2010) that was implemented in ArcGIS version 10.1 (ESRI, 2012) using custom scripts from Diltz (2015). The water balance model used monthly gridded PRISM temperature and precipitation averages (PRISM Group, 2007) for the 1971 to 2000 time period at 800-m cell size, USDA gridded SSURGO soils data (USDA, 2013) and a digital elevation model. For areas where SSURGO data were unavailable (approximately 23% of study area), we used a constant soil water holding capacity value of 100 mm which approximated the average value of the study area.

A Thornthwaite water balance model was used to calculate the following variables at monthly time steps: potential evapotranspiration (PET), actual evapotranspiration (AET), climatic water deficit (CWD), water supply (WS) and soil water balance (SWB) (Lutz *et al.*, 2010). PET represents the amount of water that would be evaporated and transpired if water were available. We follow the definitions of Stephenson (1990, 1998) and Lutz *et al.* (2010) in which AET is defined as evaporative water loss from a site covered by a hypothetical standard row crop, constrained by water availability. AET can be thought to represent net primary productivity as it considers the amount of water that is biologically useable, given the simultaneous availability of both water and energy (Stephenson, 1998). CWD is calculated as the difference between PET and AET and represents the unmet evaporative demand for water (Stephenson, 1990). Water supply is calculated as the sum of snowmelt and rain. Soil water balance is calculated as water supply minus PET plus the previous month's soil water constrained by the maximum soil water holding capacity.

To represent seasonal temperature effects, growing degree-days (GDD) was calculated as:

$$\text{GDD} = \frac{T_{\max} + T_{\min}}{2} - T_{\text{base}},$$

where T_{base} was 5.56 °C and $T_{\max}$ and $T_{\min}$ were calculated from PRISM climate data (PRISM Group, 2007). The water balance model accounted for water storage in the snowpack and soil (Lutz *et al.*, 2010) and incorporated topographic effects on solar radiation and potential evapotranspiration

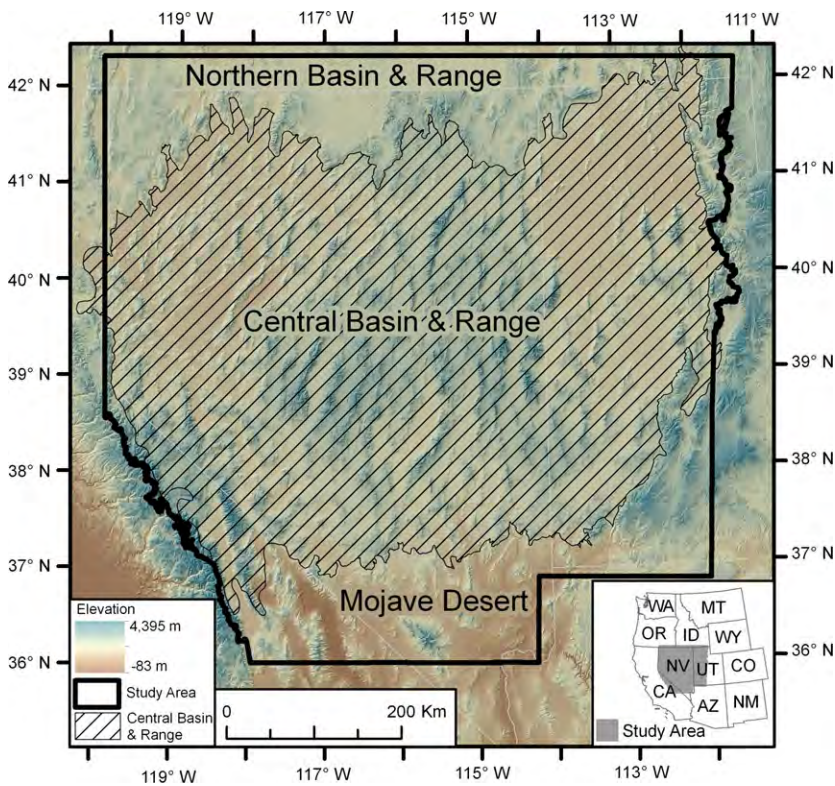


Figure 1 The study area is located in the western USA (primarily in the states of Nevada, Utah and California, 36–43° N and 111–120° W) and includes nearly all of Omernik's (1987) central basin and range province as well as portions of the northern basin and range province and the Mojave Desert. Map projection is Universal Transverse Mercator Zone 11 NAD83.

using the dimensionless heat-load index of McCune & Keon (2002) following the method of Lutz *et al.* (2010). Although the use of the heat-load index in the water balance allows for predictions at the grain of the digital elevation model, we retained the 800-m resolution for all water balance variables to compare with the bioclimatic model.

Monthly outputs of water balance (PET, AET, CWD, WS, SWB) and GDD were plotted as climographs (*sensu* Stephenson, 1998). Eighteen variables were calculated that described the magnitude, duration and timing of water balance characteristics (Fig. 2, and see Appendix S1 in Supporting Information). Rather than use all monthly variables (6 variables $\times$ 12 months = 72 variables), we used a data reduction procedure (principal components analysis) to reduce multicollinearity and determine the dimensionality of the data (see Appendix S2 for details). The following five bioclimatic variables were retained for modelling: mean annual temperature (Temp), temperature annual range (TempRang), annual precipitation (Precip), precipitation of the driest month (PrecipDryMo) and precipitation seasonality (PrecipSeas) (Table 3). The following five water deficit variables were retained for modelling: cumulative water deficit (CumLCWD), steepest rate of decline of actual evapotranspiration (DeclAET), difference between actual evapotranspiration summer low and fall peak (FallAET), positive difference in water supply between August and June (monsoonality; Ws86c) and difference between water supply and AET during the spring (WsAETSpr) (Table 4). Maps of predictor variables are shown in Appendix S2.

Statistical analysis

We used boosted regression trees (Ridgeway, 2006) to calculate separate distribution models for the 18 desert shrub species using either the bioclimatic variables or water deficit variables. The boosted regression trees method is one of the most useful techniques for presence-absence species distribution modelling, because of the good predictive performance of the technique compared to more traditional methods, such as generalized linear models and classification and regression trees, its ability to deal with interactions and nonlinear relationships, and its use of cross-validation to reduce over-fitting (Elith *et al.*, 2008). We used a tree complexity of three (roughly equivalent to the interaction order) and set the learning rate to 0.005 to ensure that a minimum of 1000 trees were generated for each species modelled.

Two methods were used to evaluate the models. First, the Southwest Regional GAP database of occurrence points was randomly partitioned into separate training ($n = 27,248$) and validation ($n = 13,624$) datasets after the removal of riparian and burned points, and the threshold-independent area under the curve (AUC) of the receiver operating characteristic (ROC) plot (Fielding & Bell, 1997) was calculated for the validation dataset (TeAUC = 'average test AUC'). Performance measures for randomly partitioned data may be inflated by spatial autocorrelation where the locations of training and validation points are close to one another (Hijmans, 2012). Therefore, we adopted an approach in which the study area was divided into four quadrants, models were

developed based on data from one quadrant and then validated using the other three quadrants as independent test data. This resulted in a total of 12 validation AUC values

Table 1 List of the shrub species used in this study.
n = number of samples in the training dataset.

Shrub code	Common name	Scientific name	<i>n</i>
ARTRW8	Wyoming sagebrush	<i>Artemisia tridentata wyomingensis</i>	2172
CHVI8	Yellow rabbitbrush	<i>Chrysothamnus viscidiflorus</i>	1787
ATCO	Shadscale	<i>Atriplex confertifolia</i>	1486
ARNO4	Black sagebrush	<i>Artemisia nova</i>	1266
ARTRV	Basin big sagebrush	<i>Artemisia tridentata tridentata</i>	1082
SAVE4	Black greasewood	<i>Sarcobatus vermiculatus</i>	850
PICRO	Bud sagebrush	<i>Picrothamnus desertorum</i>	696
SABA14	Bailey's greasewood	<i>Sarcobatus baileyi</i>	623
EPNE	Nevada ephedra	<i>Ephedra nevadensis</i>	562
KRLA2	Winterfat	<i>Krascheninnikovia lanata</i>	519
GUSA2	Broom snakeweed	<i>Gutierrezia sarothrae</i>	399
PUTR2	Antelope bitterbrush	<i>Purshia tridentata</i>	371
CELE3	Mountain mahogany	<i>Cercocarpus ledifolius</i>	327
ARAR8	Low sagebrush	<i>Artemisia arbuscula</i>	319
GRSP	Spiny hopsage	<i>Grayia spinosa</i>	303
TEGL	Littleleaf horsebrush	<i>Tetradymia glabrata</i>	246
EPVI	Mormon tea	<i>Ephedra viridis</i>	222
ATCA2	Four-wing saltbush	<i>Atriplex canescens</i>	166

Table 2 Species prevalence, mean elevation, mean soil water holding capacity, mean temperature, mean precipitation, mean growing degree-days, mean annual actual evapotranspiration and mean annual climatic water deficit for occurrence points used in this study. Prev = prevalence of species in the dataset (proportion of occurrence points occupied by the species). Elevation is in metres above sea level. SWHC = soil water holding capacity in the upper 1.5 m, Temp = average annual temperature, Precip = cumulative annual precipitation in mm, GDD = cumulative annual growing degree-days, AET = actual evapotranspiration, CWD = cumulative annual climatic water deficit. Species prevalence is based on 40,872 occurrence points.

Shrub species	Prev	Elev	SWHC	Temp	Precip	GDD	AET	CWD
ARTRW8	0.080	1776	106	8.5	554	1798	277	321
CHVI8	0.066	1869	101	8.2	582	1752	271	317
ATCO	0.054	1509	113	10.0	432	2201	223	451
ARNO4	0.046	1902	70	8.5	572	1779	273	317
ARTRV	0.039	2229	90	6.4	956	1338	325	170
EPNE	0.032	1669	68	10.9	462	2321	235	468
SAVE4	0.025	1482	168	9.7	427	2110	222	435
PICRO	0.023	1557	108	10.1	397	2161	206	463
KRLA2	0.021	1614	123	9.7	448	2105	232	427
SABA14	0.019	1519	83	10.8	352	2280	184	508
PUTR2	0.015	2029	87	7.7	784	1607	322	241
GRSP	0.014	1591	90	10.0	450	2120	231	426
GUSA2	0.012	1641	91	10.1	571	2227	284	391
ARAR8	0.012	2090	83	7.0	782	1487	310	221
CELE3	0.012	2356	88	6.0	982	1248	328	150
EPVI	0.009	1772	81	9.7	543	2098	267	383
TEGL	0.009	1553	89	10.1	437	2204	223	447
ATACA2	0.006	1501	114	11.2	382	2414	199	523

that were then averaged (Radosavljevic & Anderson, 2014; GeoAUC = 'geographically partitioned AUC').

We measured the tendency for model overfit by taking the difference between average training AUC in each of the four quadrants and average validation AUC for the validation quadrants (Radosavljevic & Anderson, 2014). Predictive maps were calculated at 800-m cell size for both types of models and all 800-m pixels were used to calculate correlation coefficients between the CWD and bioclimatic model predictions.

The relative influence of each variable was calculated using the method of Friedman (2001) as implemented using the R package GBM (Ridgeway, 2006) where each variable's influence is based on the number of times it is selected for splitting the regression tree weighted by the squared improvement to the model as the result of the split averaged over all regression trees (Elith *et al.*, 2008). The relative influence of all variables sum to 100%.

RESULTS

Climographs

The seasonal progression of site water balance varied consistently among shrub species, many of which were clearly distinguishable on the basis of their climographs (Fig. 3). The most xeric sites were occupied by black greasewood (*Sarcobatus vermiculatus*) and shadscale (*Atriplex confertifolia*) and exhibited large cumulative climatic water deficits, a high number of cumulative growing degree-days, and a small seasonal spike in water supply that barely exceeded soil water holding capacity. Black greasewood's climograph closely

Table 3 Eleven bioclimatic variables derived from the PRISM dataset. Five variables were used in the final analysis: Temp, TempRang, Precip, PrecipSeas and PcDryMo (shown in bold).

Abbreviation	Variable description
Temp	Annual mean temperature
DiurRang	Mean diurnal range (Mean of monthly (max temp – min temp))
Isotherm	Isothermality ((DirRang/TempRang) * 100)
TempSeas	Temperature seasonality (standard deviation * 100)
MaxTemp	Maximum temperature of the warmest month
MinTemp	Minimum temperature of the coldest month
TempRang	Temperature annual range (MaxTemp – MinTemp)
Precip	Annual precipitation
PcWetMo	Precipitation of the wettest month
PcDryMo	Precipitation of the driest month
PrecipSeas	Precipitation seasonality

resembled that of shadscale, but had higher variability in most water deficit variables and a less precipitous decline in AET between May and July. Black sagebrush and Wyoming big sagebrush occupied the intermediate portion of the dominant moisture gradient. Black sagebrush had more variable spring water supply in excess of PET, slightly higher variability in growing degree-days, and slightly smaller average cumulative CWD. Higher soil water holding capacity of Wyoming big sagebrush sites allowed a lower cumulative water deficit and higher cumulative AET than would be expected based on temperature and precipitation alone (Table 2). Winterfat occupied the lower to middle portion of the dominant moisture gradient with a slightly larger but more variable number of cumulative growing degree-days, a smaller water supply in excess of PET in the spring compared to Wyoming big sagebrush and black sagebrush and larger cumulative annual water deficit.

On the mesic end of the dominant moisture gradient low sagebrush and mountain big sagebrush had climographs with smaller cumulative water deficit, a smaller number of cumulative growing degree-days, and a spike in spring water supply that greatly exceeded PET. Mountain big sagebrush differed from low sagebrush in that AET was more variable among sites during the growing season.

Model performance of bioclimatic versus water deficit models

Model performance was similar for both the bioclimatic and water deficit approaches but varied greatly by species (Table 5). Average test AUC was 0.818 for the bioclimatic models (range 0.733–0.929) and 0.799 for the water deficit models (range 0.718–0.915). When evaluated against geographically independent validation areas the AUC values were, on average, 0.106 and 0.103 lower for bioclimatic and water deficit models compared to the randomly partitioned dataset. Mountain mahogany (*Cercocarpus ledifolius*) remained the best predicted species when evaluated against geographically independent validation data with an AUC of 0.865 and 0.883 for the bioclimatic and water deficit models.

Table 4 Eighteen variables that were calculated from and displayed in the water balance climographs. Five variables were retained in the final analysis: CumlCWD, DeclAET, FallAET, Ws86c and WsAETSpr (shown in bold).

Abbreviation	Variable description
DurCWD	Duration of climatic water deficit during the summer
OnsetCWD	Onset of climatic water deficit during the summer
MaxMoCWD	Month of maximum climatic water deficit
MaxCWD	Magnitude of climatic water deficit
CumlCWD	Cumulative annual climatic water deficit
DeclAET	Steepest rate of decline of actual evapotranspiration
CumIAET	Cumulative annual actual evapotranspiration
AETgdd	Cumulative annual actual evapotranspiration during the growing season
MagnSWB	Magnitude of soil water balance
FallAET	Difference between actual evapotranspiration summer low and fall peak
WsSWB	Difference between water supply and soil water balance
CumlGDD	Cumulative annual sum of growing degree-days
DurGDD	Duration of growing degree-days
SWBrec	Steepest rate of decline of soil water balance
SWBrecMo	Month of the maximum rate of soil water balance decline
Ws86c	Positive difference in water supply between August and June
WsAETSpr	Difference between water supply and AET during the spring
WsAETFall	Difference between water supply and AET during the fall

Broom snakeweed (*Gutierrezia sarothrae*), Mormon tea (*Ephedra viridis*) and four-wing saltbush showed large drops in test AUC when evaluated using the geographically independent validation data. Only three species performed better with the water deficit model compared to the bioclimatic model when evaluated using the randomly partitioned dataset and five species performed better when evaluated using the geographically partitioned dataset.

Bioclimatic and water deficit modelling approaches performed equally well when evaluated using the overfit statistic (training minus validation AUC based on the four quadrants) with nine species performing better for each of the modelling approaches (Table 5). Consistently across both approaches, models for broom snakeweed and Mormon tea were most overfit while mountain big sagebrush was least overfit (Table 5).

Correlations among modelling methods varied widely with species ranging from a low of 0.301 for broom snakeweed to a high of 0.827 for mountain big sagebrush (Table 5), with a mean value of 0.606. Most of the predicted species distribution maps appear similar regardless of whether the bioclimatic or the water deficit approach was used (Fig. 4, Appendix S3). Less overfit models yielded more similar mapped predictions as evidenced by the relationship between GeoDif and map correlation ($r = 0.514$ for bioclimatic and $r = 0.638$). In contrast, test AUC did not exhibit a strong relationship with map

correlation ($r = 0.020$ for bioclimatic and $r = 0.055$). Species in which the mapped predictions differed the most tended to occupy habitat along the Great Basin–Mojave Desert transitional zone and the mapped differences clearly exhibited a geographically clustered pattern (Fig. 4).

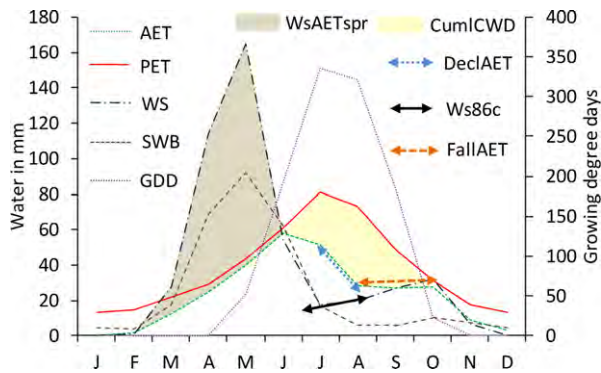


Figure 2 Climograph for a typical site dominated by low sagebrush showing the five water deficit variables used for predictive modelling in this study. Variables were calculated as area under the curve (CumICWD, cumulative annual climatic water deficit; WsAETspr, difference between water supply and actual evapotranspiration in the spring), and recession rate (DeclAET, steepest rate of decline of actual evapotranspiration; FallAET, difference between actual evapotranspiration summer low and fall peak; Ws86c, positive difference in water supply between August and June) based on monthly CWD variables (PET, potential evapotranspiration; AET, actual evapotranspiration; WS, water supply; SWB, soil water balance; GDD, growing degree-days).

Relationship between shrub species and ecological gradients

For models using the water deficit approach, the cumulative water deficit variable was the most important predictor variable with a mean relative influence of 31% compared to the next highest variable (WsAETspr), with a mean relative influence of 22% (Fig. 5). Cumulative water deficit was the most important predictor variable for 10 out of 18 species with spring water supply (WsAETspr) the most predictive for five species, followed by monsoonality (Ws86c) for two species, and fall AET for one species. In contrast, the bioclimatic models showed less agreement in which variable was selected as the most important. The top variable, temperature, had a mean relative influence of 24% which was only slightly different from the next highest, precipitation, which had a mean importance of 23%. The number of species in which temperature was the best predictor was five, followed by precipitation (five), temperature range (four), precipitation seasonality (two), and precipitation in the driest month (two).

Four of the five species that had spring water supply as their most important water deficit predictor variable were the montane species (mountain mahogany, mountain big sagebrush, low sagebrush and antelope bitterbrush; see Table 2 for their average elevation) with black greasewood being the exception (Appendix S3). Fall AET was the most predictive water deficit variable for winterfat. Broom snakeweed and Mormon tea were best predicted by monsoonality, which is reflected in their modelled distribution being

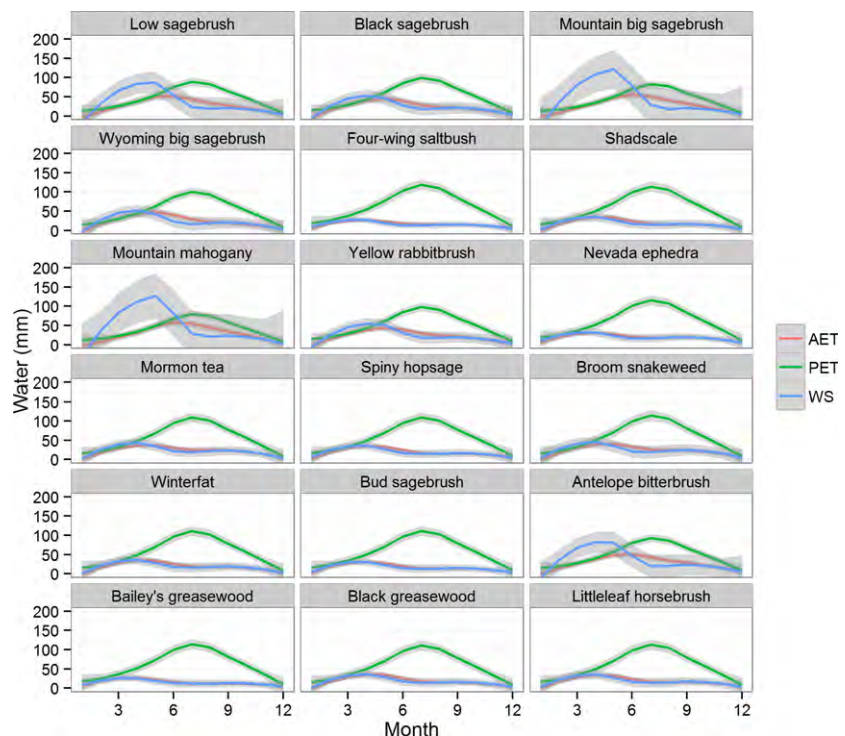


Figure 3 Climographs based on water balance variables for all study species. Grey shading indicates 95% confidence intervals using all sites. AET, actual evapotranspiration; PET, potential evapotranspiration; WS, water supply (rain + snowmelt).

Table 5 Model performance measures for all 18 shrub species using models derived with bioclimatic variables (BIO) and climatic water deficit variables (CWD). TestAUC = AUC of the validation data, GeoAUC = average test AUC of the geographically partitioned data when models from one quadrant were applied to the other three quadrants, GeoDif = difference between the training and test AUC using the geographically partitioned data (higher values are more likely to indicate overfitting) and Corr = map-wide correlation between bioclimatic and water deficit using all 800-m pixels from the predictive map.

Shrub species	TestAUC		GeoAUC		GeoDif		Corr
	BIO	CWD	BIO	CWD	BIO	CWD	
ARTRW8	0.790	0.783	0.715	0.740	0.135	0.093	0.811
CHVI8	0.733	0.718	0.662	0.617	0.127	0.160	0.733
ATCO	0.795	0.773	0.742	0.711	0.098	0.122	0.700
ARNO4	0.802	0.777	0.723	0.703	0.154	0.150	0.668
ARTRV	0.875	0.857	0.853	0.842	0.052	0.049	0.827
EPNE	0.875	0.805	0.708	0.702	0.194	0.189	0.516
SAVE4	0.803	0.784	0.784	0.675	0.101	0.190	0.541
PICRO	0.829	0.808	0.741	0.735	0.150	0.145	0.782
KRLA2	0.828	0.771	0.691	0.687	0.220	0.209	0.592
SABA14	0.789	0.894	0.743	0.780	0.183	0.139	0.733
PUTR2	0.913	0.810	0.772	0.750	0.153	0.151	0.397
GRSP	0.854	0.754	0.665	0.636	0.245	0.259	0.625
GUSA2	0.827	0.796	0.595	0.547	0.312	0.345	0.301
ARAR8	0.824	0.826	0.731	0.698	0.198	0.222	0.575
CELE3	0.929	0.915	0.865	0.883	0.092	0.068	0.688
EPVI	0.834	0.774	0.605	0.533	0.303	0.362	0.366
TEGL	0.778	0.790	0.679	0.687	0.217	0.227	0.587
ATCA2	0.744	0.739	0.641	0.603	0.299	0.329	0.473

centred in the south-east monsoonal portion of the study area (Appendix S3).

Similarly, three of the five species that had precipitation as their most predictive bioclimatic variable were montane species (mountain mahogany, mountain big sagebrush and low sagebrush) with bud sagebrush and four-wing saltbush being exceptions (Appendix S3). Black greasewood, shadscale, winterfat and Mormon tea were best predicted by annual temperature range. Broom snakeweed and littleleaf horsebrush (*Tetradymia glabrata*) were best predicted by precipitation seasonality, and yellow rabbitbrush (*Chrysothamnus viscidiflorus*) and black sagebrush were best predicted by precipitation in the driest month.

DISCUSSION

Does the water deficit approach yield better species distribution models?

Correlative species distribution models have been criticized for lacking mechanistic explanations and relying upon variables that are only indirectly related to plant resources, such as water, light, and nutrients (Guisan & Zimmermann, 2000). Inferences regarding controls on species distributions

may be challenging to verify without recourse to experimentation over such broad spatial scales, and can be difficult to extrapolate to future climates. Species distribution models have shown limited capacity to predict to new geographical areas (Randin *et al.*, 2006), and the lack of geographical transferability may be due to overfit models (Heikkinen *et al.*, 2012) or environmental variables that correlate with, but do not functionally influence, species distributions (Dormann, 2007). Therefore, we expected that SDMs using water deficit variables, which we considered to represent more functional constraints upon species distribution, would prove more generalizable.

Contrary to expectations, SDMs using variables constructed from a water deficit approach performed similarly to SDMs using standard bioclimatic variables for predicting arid-land shrub species distributions. Even when evaluating low-elevation shrub species that are most water-limited, species differences in predictive ability of the two modelling approaches proved idiosyncratic (Table 5). Previous research focused on tree species showed that water deficit approaches resulted in improved predictions (Piedallu *et al.*, 2013). We ascribe the lack of improvement in model performance to the ability of the bioclimatic variables to capture unique aspects of the climate that the water deficit modelling did not capture. In particular, temperature range and precipitation seasonality appear to capture aspects of climate that are not modelled very well by the water deficit approach. The water deficit variables describing fall AET, rate of decline of AET (DeclAET) and monsoonality (Ws86c) explained less of the deviance compared to temperature range and precipitation seasonality (Fig. 5).

Despite the lack of improved performance in model fit and spatial transferability, the water deficit modelling approach offers advantages over the bioclimatic approach in that its predictor variables may be more interpretable than standard bioclimatic variables, because they are more functionally relevant and distinguish shrub species according to key features of the seasonal water balance (Fig. 3). Interpretation of simple climate variables, such as mean annual temperature and precipitation, can be difficult because both variables operate in tandem to influence site water balance. Thus, use of predictor variables that serve as spatial proxies for actual water resource limitation of arid-land shrub species should result in improved capacity to explain changes in species distributions in a warming environment (Franklin, 1995; Austin, 2002). However, the water balance approach uses additional information on soils and topography that may impart some additional error.

Dominant environmental influences on shrub species distributions

Cumulative CWD was consistently selected as the most influential variable in our water deficit models. Other authors have primarily used cumulative annual CWD and cumulative annual AET for modelling plant distributions (Stephenson,

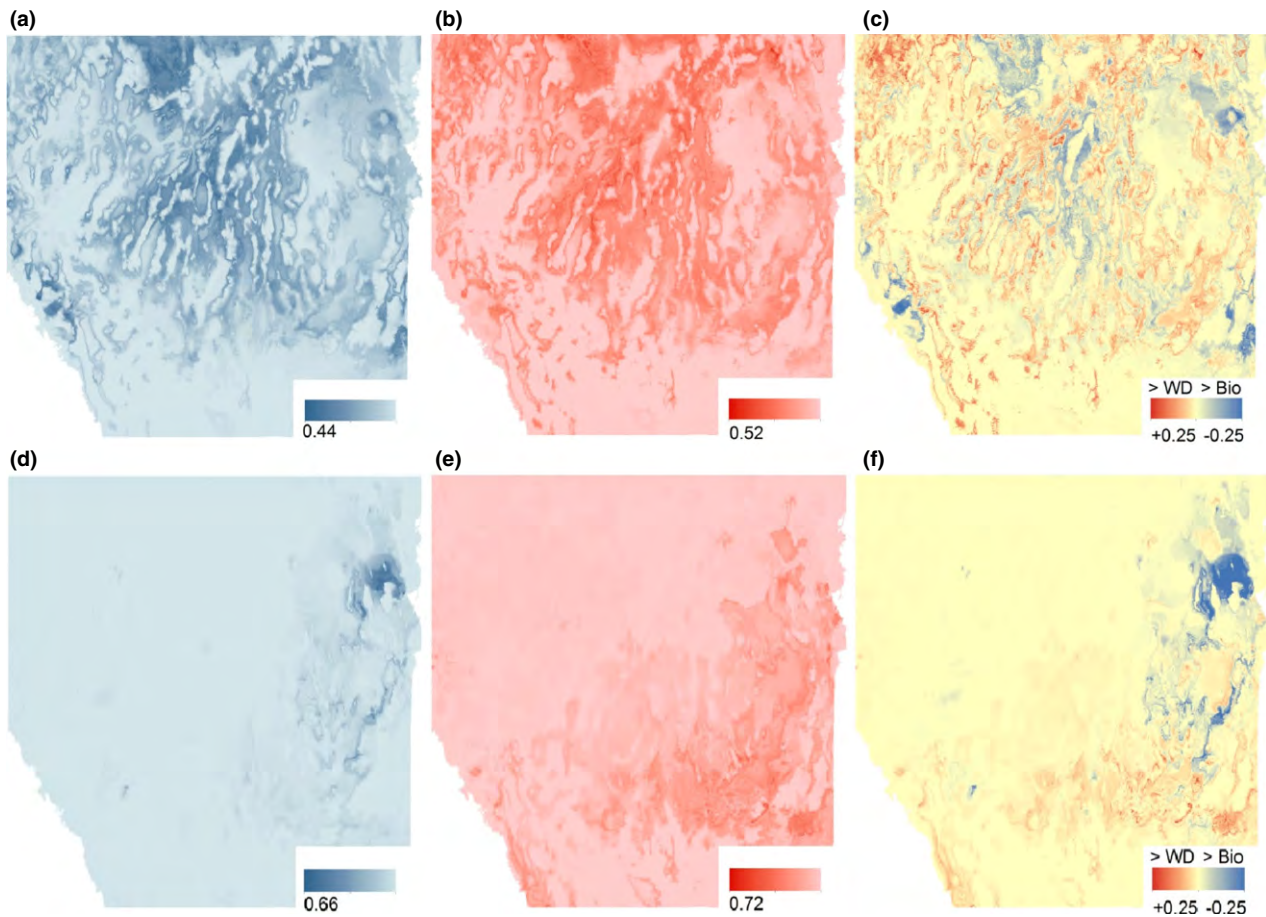


Figure 4 Maps showing predictive models for Wyoming big sagebrush and broom snakeweed using bioclimatic variables (a and d) and water deficit variables (b and e) and the difference between water deficit and bioclimatic predictions (c and f). Reds indicate that the climatic water deficit model predicted higher values of occurrence probability and blues indicate that the bioclimatic model predicted higher values of occurrence probability.

1990, 1998; Lutz *et al.*, 2010; Piedallu *et al.*, 2013). We extended this previous approach by generating and further analysing 18 additional variables derived from the CWD climograph, allowing a more synthetic description of how the shape, rate and timing of the water balance change through the year. We found several of these additional variables to be important for predicting distribution of important shrub types in the Great Basin, including difference between water supply and actual evapotranspiration in spring (WsAETSpr), difference between the minimum value of AET in summer and the maximum value of AET in fall (FallAET), difference between water supply in August and June (Ws86c), and the steepest rate of decline in AET (DeclAET). In other systems a wider range or different set of water deficit variables may be useful for describing differences among plant communities.

Relationships between water deficit variables and shrub species distribution are consistent with results from detailed ecohydrological modelling studies (Schlaepfer *et al.*, 2012a) that have linked sagebrush species occurrence to environments where spring recharge is followed by a dry period (i.e.

high WsAETSpr, high DeclAET). The spring water supply variable (WsAETSpr) is particularly important for modelling the distribution of the more montane shrub species that grow in areas where a large snowpack results in a large surplus of soil water during the wet season followed by drying during the summer growing season.

Previous studies have used water balance approaches derived from CWD modelling to predict tree species distributions (e.g. Lutz *et al.*, 2010; Crimmins *et al.*, 2011). Mechanistic process models have also been developed to quantify the ecohydrological niche of sagebrush vegetation (Schlaepfer *et al.*, 2012a), and to create species distribution models for big sagebrush (Schlaepfer *et al.*, 2012b). Our results further support the efficacy of water balance approaches, which combine information on precipitation inputs, temperature-related evapotranspiration outputs, and water holding capacity of the soil, for defining the ecohydrological niche of vegetation in water-limited landscapes. However, the ability of bioclimatic models to predict as accurately or better than water deficit models suggests that temperature has important influences on shrub species distribution that may not relate

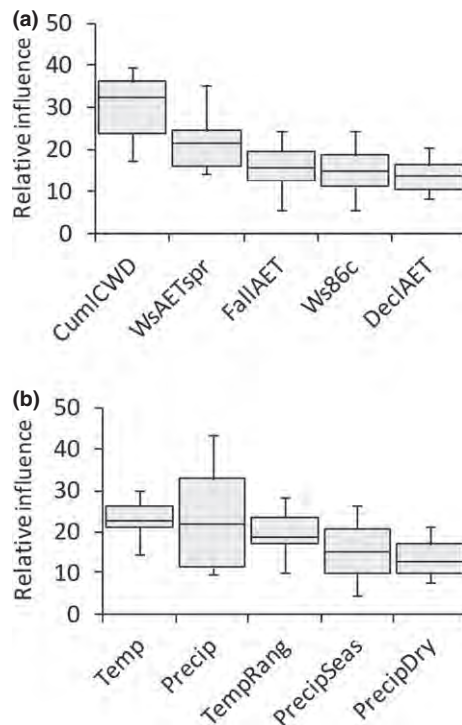


Figure 5 Boxplots showing the relative influence of each environmental predictor across all study species for models calculated from (a) bioclimatic variables and (b) water deficit variables. Horizontal lines in boxes show medians and boxes show upper and lower quartiles, with vertical lines showing minimum and maximum values.

to the water balance, such as effects on seed germination and seedling establishment (Meyer & Monsen, 1992), and eco-physiological processes including photosynthesis and respiration (e.g. Ryan, 1991; Ryel *et al.*, 2010).

Temperature effects may operate synergistically with precipitation seasonality, in that rainfall occurring during the growing season may result in optimal conditions for root and shoot growth (Noy-Meir, 1973). Improved species distribution models for arid-land shrubs could be developed by including a soil temperature component in addition to a water balance component or through the inclusion of daily gridded meteorological data to account for pulsed fluxes of water and temperature (Abatzoglou, 2013).

Implications for modelling vegetation response to climate change

Modelling vegetation response to climate change requires incorporation of the proximal mechanisms that control plant establishment, growth and survival. Climatic correlates that do not represent true causal mechanisms may yield high correlations using current climate data, yet such empirical relationships may not hold up in the future (Davis *et al.*, 1998; Dormann, 2007). The water deficit modelling approach integrates spatially distributed information on seasonally varying precipitation and temperature with high-resolution topo-

graphic and soil data, to construct a synthetic measure of functional water balance for plants. Our study shows that spatial variables derived from water deficit models result in predictions comparable to commonly used bioclimatic variables, when used to model distributions of 18 shrub species across the entire, 451,112 km² Great Basin region. We suggest that our SDM modelling approach would prove more robust to modelling range shifts given climate change scenarios than models using standard bioclimatic variables. Although our approach remains correlative in the sense of Kearney & Porter (2009), our study highlights the promise of linking existing GIS modelling techniques with commonly available spatial data for creating SDMs that use direct variables of primary importance to vegetation distribution in arid and semi-arid landscapes. Future research should develop functionally relevant, spatial variables for other direct influences on plant species distribution, such as energy budget, temperature or nutrient availability (e.g. Coudun *et al.*, 2006), and combine these with water deficit variables to derive functionally relevant, generalizable SDMs for modelling climate change scenarios.

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SUPPORTING INFORMATION

Additional Supporting Information may be found in the online version of this article:

Appendix S1 Derivation of water deficit predictors and maps showing the variables used in this study.

Appendix S2 Results of the principal components analysis for bioclimatic and water deficit variables.

Appendix S3 Predictive maps and variable influence for each species distribution model.

BIO SKETCH

Thomas Dilts is a research scientist in the Great Basin Landscape Ecology Lab at the University of Nevada Reno. His interests are in the application of geographic information systems and remote sensing techniques for understanding vegetation dynamics and for conservation planning.

All co-authors have interests in vegetation dynamics, climate change, and ecosystem resilience and resistance to disturbance.

Author contributions: T.E.D., P.J.W. and J.C.C. conceived the ideas; C.M.D. and T.E.D. analysed the data; and T.E.D., P.J.W. and J.C.C. led the writing.

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SUPPORTING INFORMATION

Functionally relevant climate variables for arid lands: using climatic water deficit modeling to predict distribution of desert shrub vegetation

T.E. Dilts, P. J. Weisberg, C. M. Dencker, and J. C. Chambers

Appendix S1 Derivation of the water deficit predictive variables used in this study and maps showing the bioclimatic and water deficit variables used in this study.

Methods

Monthly rasters of potential evapotranspiration, actual evapotranspiration, climatic water deficit, water supply, and soil water balance were calculated from a Thornthwaite water balance following the approach of Lutz *et al.* (2010) using the custom scripts (Dilts, 2015) executed in ArcGIS version 10.1 software (Esri, 2011) and assembled to create a water balance climograph. Growing degree days were included in the climograph to assess the seasonal balance of water as it compares to the growing season. All climate variables were derived from the 1970 – 2001 PRISM normals (PRISM, 2007).

Below are descriptions of how each of the five variables used in the analysis were calculated in order of their average influence on the final models of all species:

(1) Cumulative annual climatic water deficit (CumlCWD) – Sum of the difference between PET and AET for all months of the year. This can also be thought of as the area between the PET and AET curves on the climograph.

(2) Difference between water supply and AET during the spring (WsAETspr) – In snowmelt systems a large pulse in the available water supply occurs during the spring months. This variable is calculated by subtracting AET from water supply for all spring months in which water

supply exceeds AET. Water supply in excess of AET is considered to be runoff or percolation, however, the model used in this study does not consider horizontal movement of water. In the Great Basin of the USA larger values are associated with areas that have large snowpacks.

(3) Difference between actual evapotranspiration summer low and fall peak (FallAET) – In the autumn there can be an increase in AET associated with slightly increased precipitation coupled with lower PET rates. In other areas there is little increase if there is very little precipitation in the autumn. This variable is calculated as the difference in AET between autumn maximum and the summer minimum.

(4) Positive difference in water supply between August and June (monsoonality Ws86c) – Some areas of the western USA receive late summer rains as a result of the North American monsoon. This variable is calculated by taking the difference in water supply between August and June and only including the positive values. Larger values indicate an increase in precipitation over the summer usually as a result of rains from the North American monsoon.

(5) Steepest rate of decline of actual evapotranspiration (DeclAET) – The largest month-to-month drop in AET and is calculated as the difference between AET in the higher month and AET in the lower month. This typically occurs during the summer when the soil water is being depleted rapidly. In the Great Basin of the USA the largest drops occur in areas of intermediate elevation (1425 – 1835-m), because those areas have larger amounts of soil water compared to lower elevations coupled with high PET values during the summer. Furthermore, many of the areas with the highest DeclAET are located on fine-textured soils where the available water capacity is high.

The following additional thirteen variables were calculated but not included in the final analysis due to issues of multicollinearity. These variables are described below:

(6) Duration of climatic water deficit (DurCWD) – The number of months in which PET exceeds AET.

(7) Onset of climatic water deficit (OnsetCWD) – The first month of the year in which PET exceeds AET.

(8) Month of maximum climatic water deficit (MaxMoCWD) – The month with the largest difference between PET and AET.

(9) Magnitude of climatic water deficit (MaxCWD) – The maximum difference in PET and AET in millimetres. Magnitude of climatic water deficit is always calculated for the month of maximum climatic water deficit.

(10) Cumulative annual actual evapotranspiration (CumlAET) – Sum of all monthly AET values. This can also be thought of as the area under the AET curve on the climograph. In the Great Basin of the USA the highest values occurred at elevations 2,000 to 2,500-m where snowpacks were large yet the growing season was longer than on the mountain tops.

(11) Cumulative annual actual evapotranspiration during the growing season (AETgdd) – Sum of all monthly AET values when GDD is greater than zero.

(12) Magnitude of soil water balance (MagnSWB) – Calculated as the difference between the spring maximum and summer minimum soil water balance.

(13) Difference between water supply to soil water balance (WsSWB) – The sum of the difference between water supply and soil water balance for all spring months in which water supply exceeds soil water balance.

(14) Cumulative annual sum of growing degree days (CumlGDD) – Sum of growing degree days for all months in the year.

(15) Duration of growing degree days (DurGDD) – Number of months in which the growing degree days are greater than zero.

(16) Steepest rate of decline of soil water balance (SWBrec) – The largest month-to-month drop in soil water balance and is calculated as the difference between soil water balance in the higher month and soil water balance in the lower month.

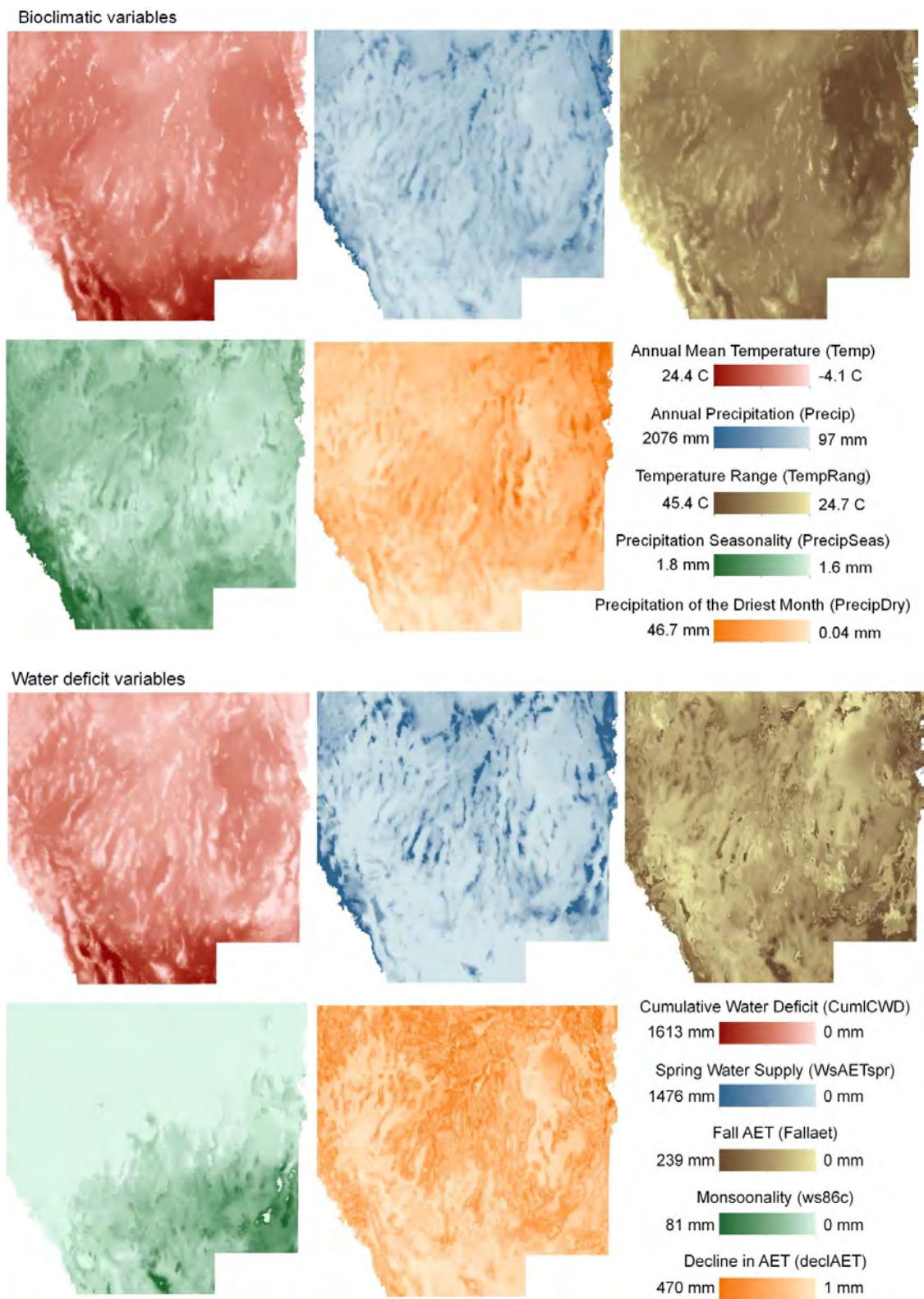
(17) Month of the maximum rate of soil water balance decline (SWBrecMo) – The month with the largest drop in soil water balance.

(18) Difference between water supply and AET during the fall - This variable is calculated by subtracting AET from water supply for all autumn months in which water supply exceeds AET.

Table S1 Mean, standard deviation, minimum, and maximum values for all raster cells (800-m) for the eighteen water balance variables examined in this study. Variables highlighted in bold were used in the final analysis. All units are in millimetres except for the following variables: DurCWD, OnsetCWD, MaxMoCWD, DurGDD, and SWBrecMo. DurCWD and DurGDD are measured in number of months. OnsetCWD, MaxMoCWD, and SWBrecMo identify a particular month in which a maximum value occurs and are numbered 1 = January through 12 = December.

Abbreviation	Mean	St. Deviation	Minimum	Maximum
CumlcWD	378.96	212.40	0.00	1613.79
WsAETSpr	90.74	126.56	0.00	1476.30
FallAET	12.51	6.42	0.00	238.96
Ws86c	5.62	8.58	0.00	81.13
DeclAET	32.07	15.91	1.46	469.74
DurCWD	6.45	1.75	0	9
OnsetCWD	5.38	2.13	3	10
MaxMoCWD	7.26	0.53	6	10
MaxCWD	95.65	35.33	0.00	529.86
CumIAET	265.16	77.02	29.04	642.27
AETgdd	204.99	66.88	0.00	551.60
MagnSWB	47.53	40.06	0.00	301.50
WsSWB	206.78	100.55	10.08	1643.82
CumIGDD	2067.55	813.89	0.00	6143.81
DurGDD	7.15	1.52	0	12
SWBrec	29.98	21.20	0.00	237.08
SWBrecMo	5.05	1.27	2	12
WsAETFal	2.87	10.79	0.00	833.61

Figure S1 Maps showing the bioclimatic and water deficit predictive variables used in this study.



SUPPORTING INFORMATION

Functionally relevant climate variables for arid lands: using climatic water deficit modeling to predict distribution of desert shrub vegetation

T. E. Dilts, P. J. Weisberg, C. M. Dencker, and J. C. Chambers

Appendix S2 Results of the principal components analysis for bioclimatic and water deficit variables.

Methods

Principal components analysis (PCA) was performed on the bioclimatic variables and the water deficit variables using the training dataset derived from the Southwest ReGAP database (USDI, 2004a) using PcOrd software (McCune and Medford, 2011) in order to determine the main axes of variation in the data. The PCA was calculated based on the correlation matrix of bioclimatic variable and water deficit variables separately and 999 randomizations were run to determine axis significance. Species means were overlaid on the PCA biplots to determine how the species clustered in environmental space.

Results

For the bioclimatic variables the first four axes were significant explaining 94% of the variance in the data. The first axis accounted for 51% of the variance explained while the second, third, and fourth axes accounted for 19%, 15%, and 9% of the variance, respectively. The first axis was positively correlated with annual precipitation ($r=0.92$), precipitation of the wettest month ($r=0.91$), and negatively correlated maximum temperature ($r=-0.82$) while the second axis was positively correlated with minimum temperature ($r=0.87$), and negatively correlated with

annual temperature range ($r=-0.62$) and temperature seasonality ($r=-0.45$) (Fig. S1a). The third axis was positively correlated temperature seasonality ($r=0.62$) and negatively correlated with isothermality ($r=-0.82$) and precipitation seasonality ($r=-0.47$) (Fig. S1b). The fourth axis was positively precipitation seasonality ($r=-0.75$).

For the water deficit variables the first three axes were significant explaining 84% of the variance. The first axis accounted for 64% of the variance explained while the second and third axes each accounted for 10%. Cumulative water deficit ($r=0.94$), duration of water deficit ($r=0.93$), and maximum water deficit ($r=0.92$) were all highly positively correlated with axis 1 while the magnitude of soil water balance ($r=-0.91$) was most negatively correlated with axis 1 (Fig. S1c). The second axis was positively correlated with the increase in fall AET ($r=0.75$) and monsoonality (ws96c; $r=0.72$) (Fig. S1c). The third axis was positively correlated with the rate of decline in AET during the summer ($r=0.58$) and AET during the growing season ($r=0.48$) and negatively correlated excess water supply in the fall ($r=-0.48$) and month of maximum water deficit ($r=-0.41$) (Fig. S1d).

The species means on the PCA biplot indicate that the four xeric montane species (low sagebrush, bitterbrush, mountain big sagebrush, and mountain mahogany) were separated from all of the other species along axis 1 of both the bioclimatic and water deficit variables (Fig. S2). Axes 2 and 3 for the bioclimatic PCA did a better job of separating the species from one another compared axes 2 and 3 of the water deficit PCA. In contrast, the water deficit PCA yielded four clusters of species representing montane xeric species, sagebrush steppe species (Wyoming big sagebrush, yellow rabbitbrush, and black sagebrush), salt desert scrub species (shadscale, greasewood, bud sagebrush, winterfat, littleleaf horsebrush, and spiny hopsage) and Great Basin-Mojave transitional species (broom snakeweed, mormon tea, green ephedra, and four wing

saltbush) (Fig. S2c). The latter two groups were separated by axis 2 of the water deficit PCA while axis 2 (Fig. S2a) of the bioclimatic PCA separated Bailey's greasewood, four-wing saltbush, green ephedra, mormon tea, and spiny hopsage from the remaining salt desert scrub species. Axis 3 of both the bioclimatic and water deficit PCAs (Figures S2b and S2d) separated sticky snakeweed and green ephedra from the remaining salt desert scrub and sagebrush steppe species while axis 3 of the bioclimatic PCA (Fig. S2d) additionally separated littleleaf horsebrush and shadscale from the remaining salt desert scrub species.

Correlations among water deficit and bioclimatic variables showed an average correlation of 0.45 when ignoring the sign of the correlation. Isothermality, temperature seasonality, temperature range, and precipitation seasonality were relatively uncorrelated with water deficit variables and had correlations with water deficit variables -0.5 to 0.5 while diurnal range had correlations -0.6 to 0.6 (Fig. S3). The water deficit variables that were uncorrelated with bioclimatic variables were rate of decline in AET (declAET), increase in fall AET (FallAET), and monsoonality (Ws86c) all having correlations between -0.5 to 0.5.

Fig. S1 Principal components biplots showing the variable loadings for (a) bioclimatic variables with axis 1 and 2, (b) bioclimatic variables with axis 1 and 3, (c) water deficit variables with axis 1 and 2, (d) water deficit variables with axis 1 and 3.

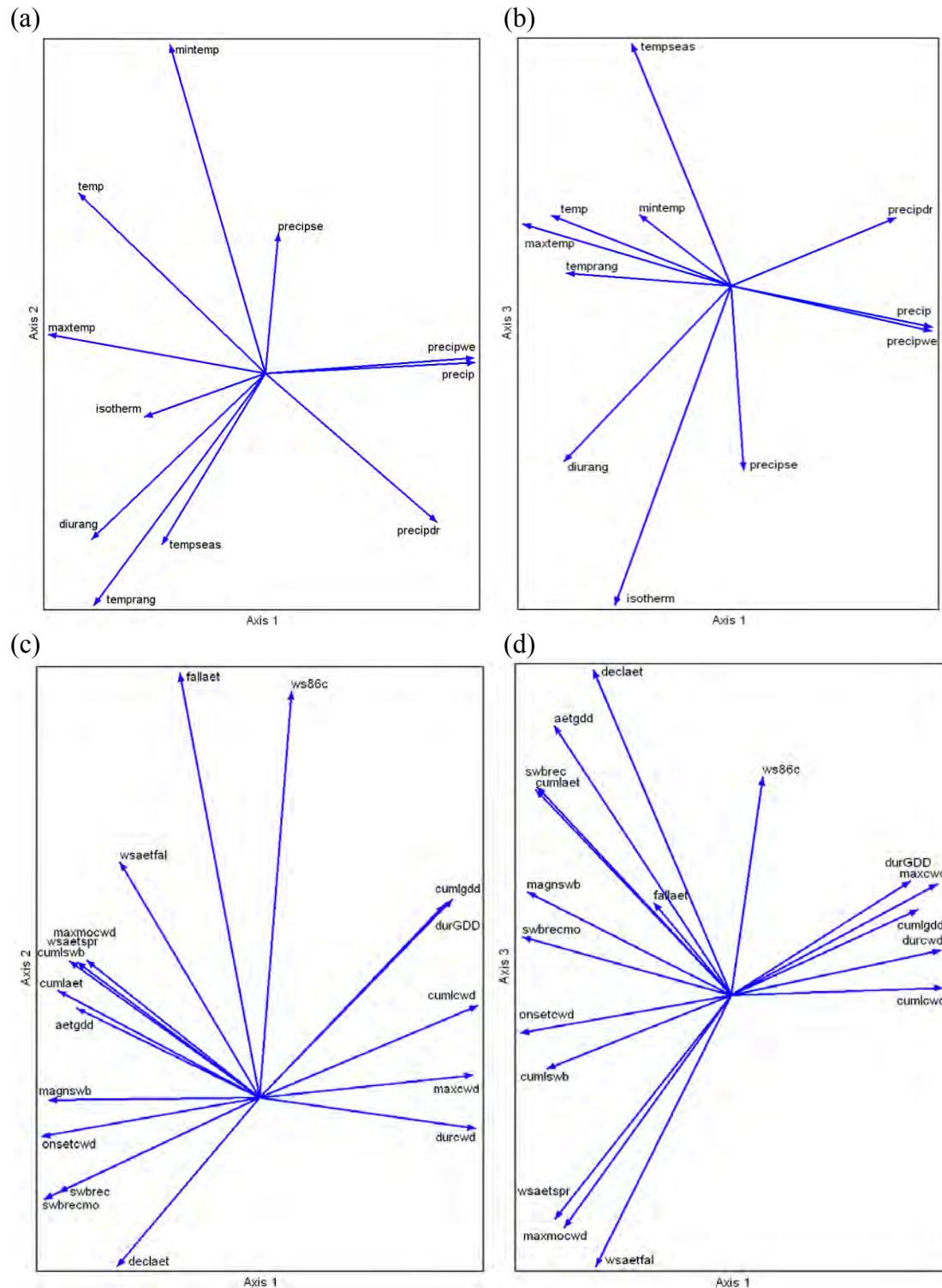
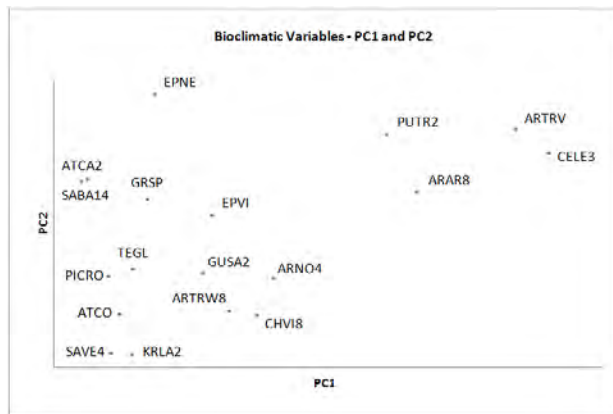
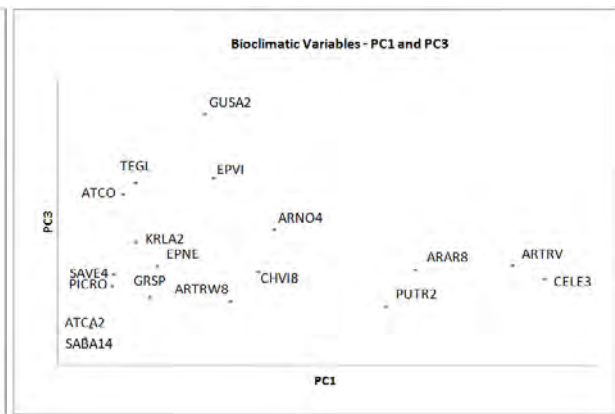


Fig. S2 Principal components biplots showing the species means (blue dots) for (a) bioclimatic variables with axis 1 and 2, (b) bioclimatic variables with axis 1 and 3, (c) water deficit variables with axis 1 and 2, (b) water deficit variables with axis 1 and 3.

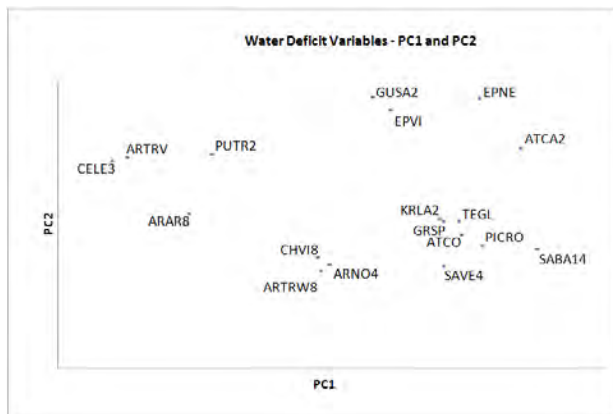
(a)



(b)



(c)



(d)

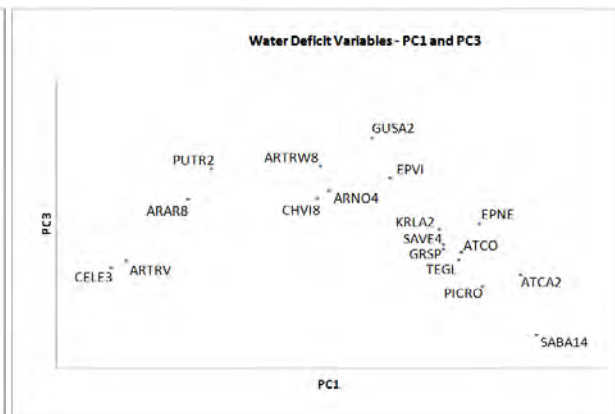


Fig. S3 Pearson product-moment correlation coefficient (r) between bioclimatic and water deficit variables. Bioclimatic variables are shown as columns and water deficit variables as rows. The colour of the cell indicates the strength of the correlation (red = ≥ 0.9 or ≤ -0.9 , orange = 0.8 to 0.9 or -0.9 to -0.8, green = 0.7 to 0.8 or -0.8 to -0.7, blue = 0.6 to 0.7 or -0.7 to -0.6, white = -0.6 to 0.6). Minus symbols indicate a negative relationship and the lack of a minus symbol indicates a positive relationship.

	aetgdd	cumlaet	cumlcwd	cumlgdd	Cumlswb	declaet	durcwd	durGDD	fallaet	magnswb	maxcwd	maxmocwd	onsetcwd	swbrec	swbreemo	ws86c	wsaetfal	wsaetspr
temp	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
diurang	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
isotherm	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
tempseas	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
maxtemp	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
mintemp	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
temprang	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
precip	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
precipwet	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
precipdry	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
precipseas	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-

Reference

McCune, B. & Mefford, M.J. (2011) *PC-ORD. Multivariate Analysis of Ecological Data. Version 6*. MjM Software, Gleneden Beach, Oregon, U.S.A.

U.S. Department of the Interior. (2004a) *U.S. Geological Survey National Gap Analysis Program, Southwest Regional Gap Analysis Project Field Sample Database. Version 1.1*. RS/GIS Laboratory, College of Natural Resources, Utah State University.

SUPPORTING INFORMATION

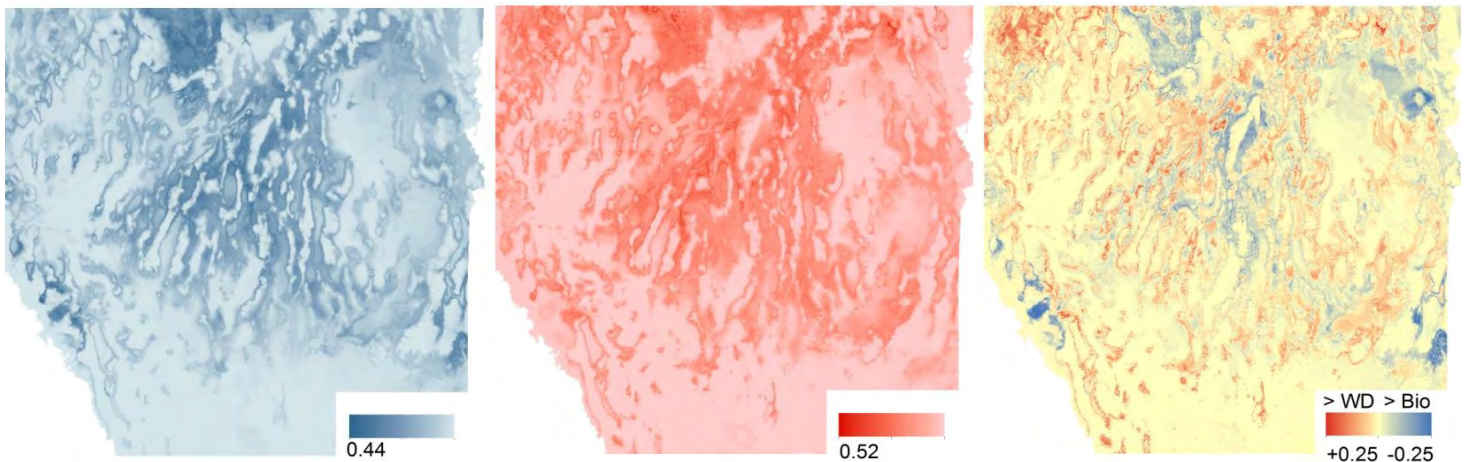
Functionally relevant climate variables for arid lands: using climatic water deficit modeling to predict distribution of desert shrub vegetation

T.E. Dilts, P. J. Weisberg, C. M. Dencker, and J. C. Chambers

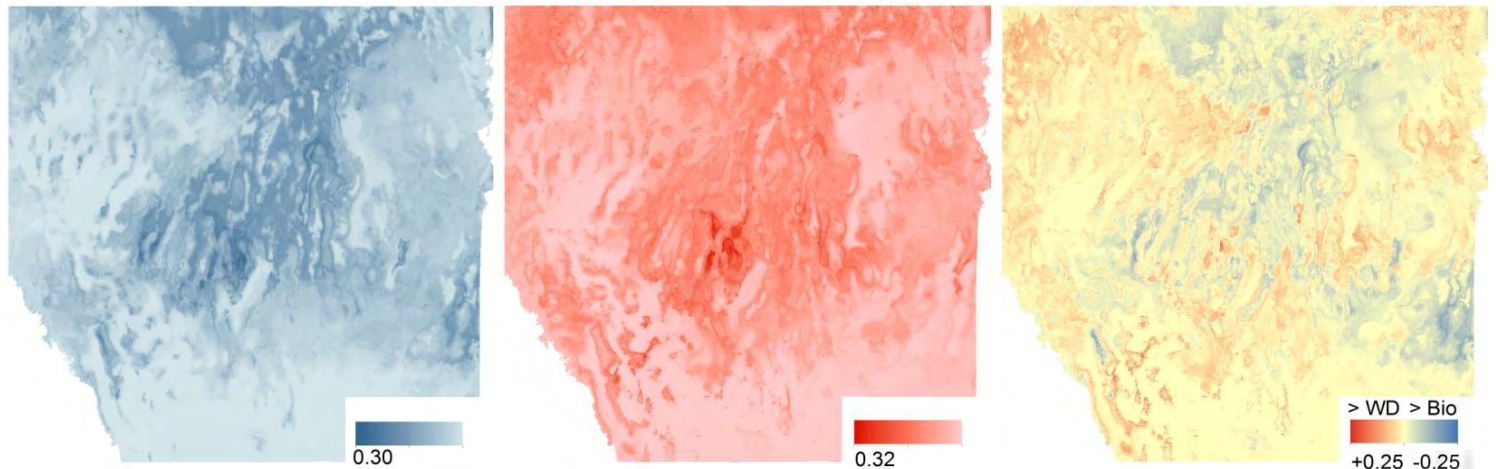
Appendix S3 Predictive maps and variable influence for each species distribution model.

Figure S1 Predictive maps for each species using models generated from bioclimatic variables (left), water deficit variables (centre), and the difference between water deficit and bioclimatic predictions (right). Colours in the maps represent linear scaling.

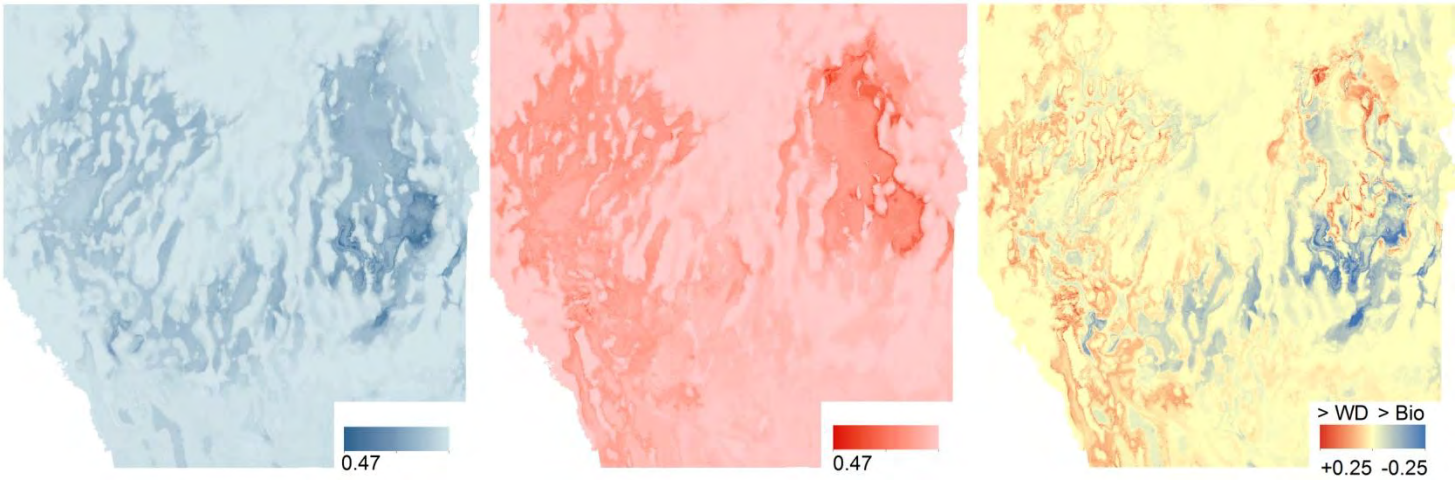
Wyoming big sagebrush (*Artemisia tridentata wyomingensis*)



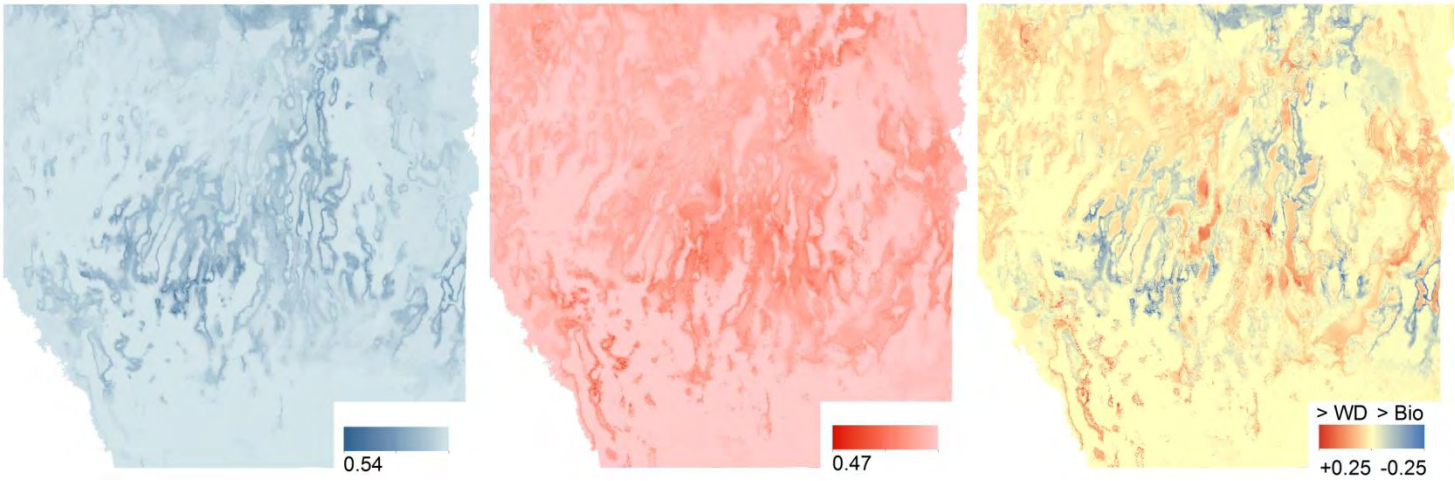
Yellow rabbitbrush (*Chrysothamnus viscidiflorus*)



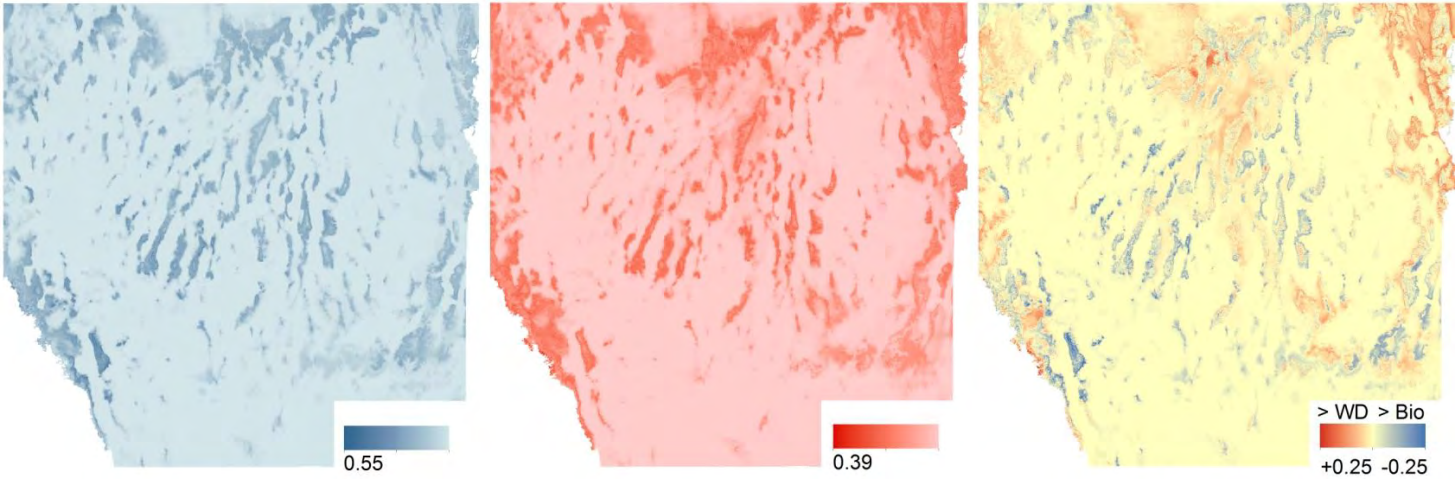
Shadscale (*Atriplex confertifolia*)



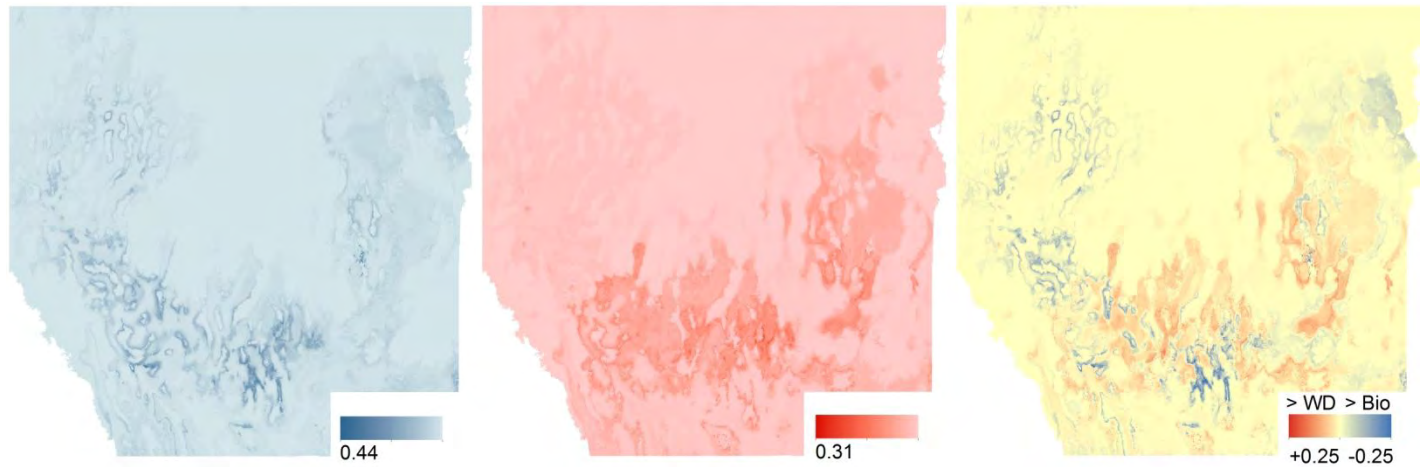
Black sagebrush (*Artemisia nova*)



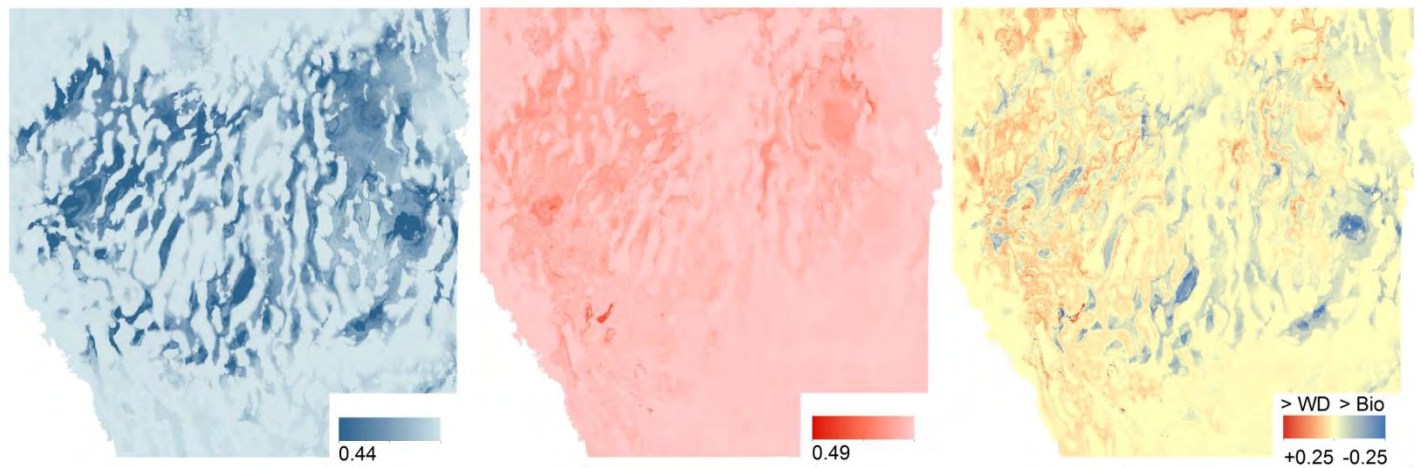
Mountain big sagebrush (*Artemisia tridentata vaseyana*)



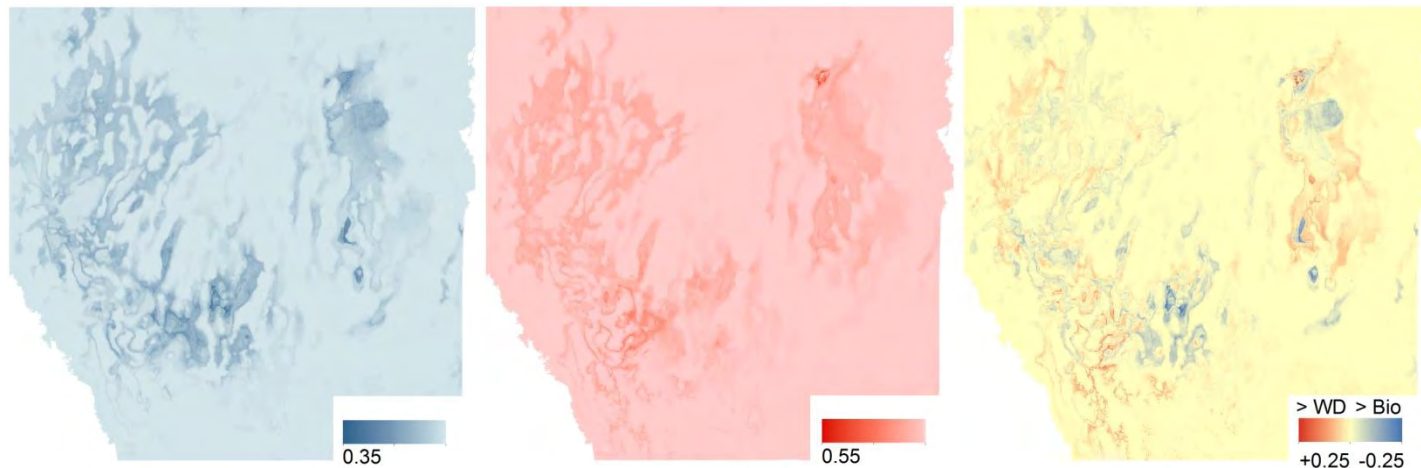
Nevada ephedra (*Ephedra nevadensis*)



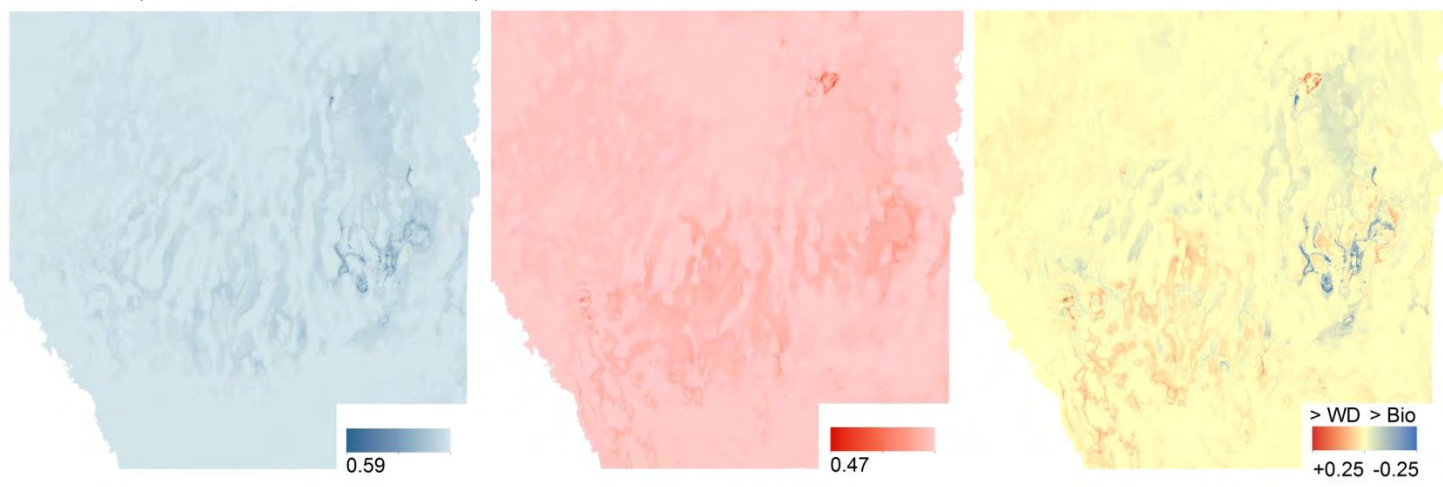
Black greasewood (*Sarcobatus vermiculatus*)



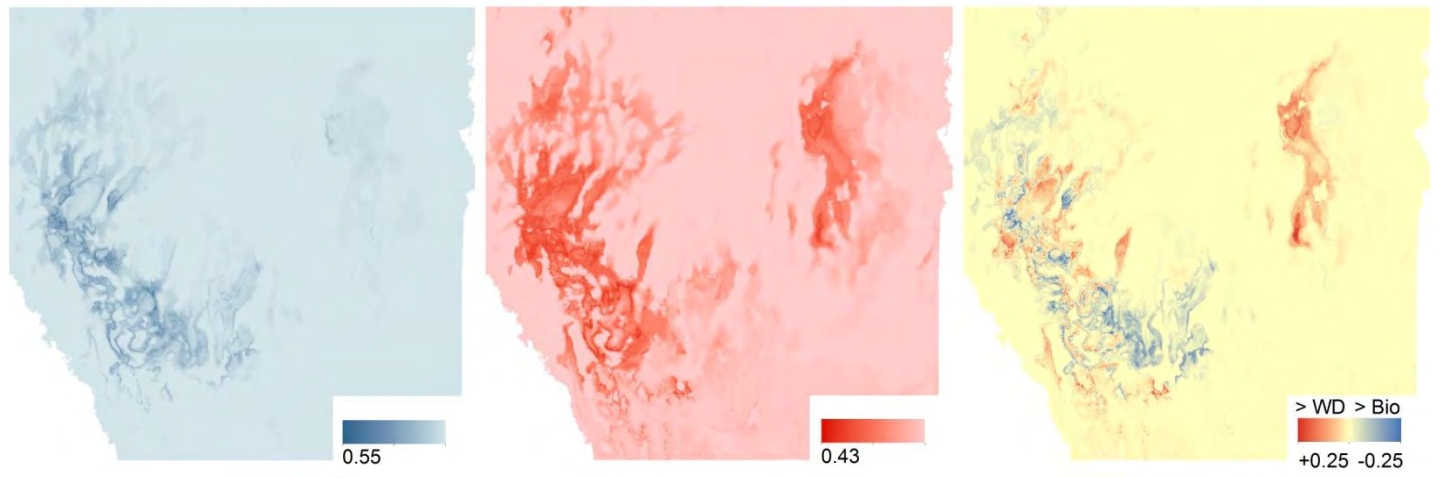
Bud sagebrush (*Picrothamnus desertorum*)



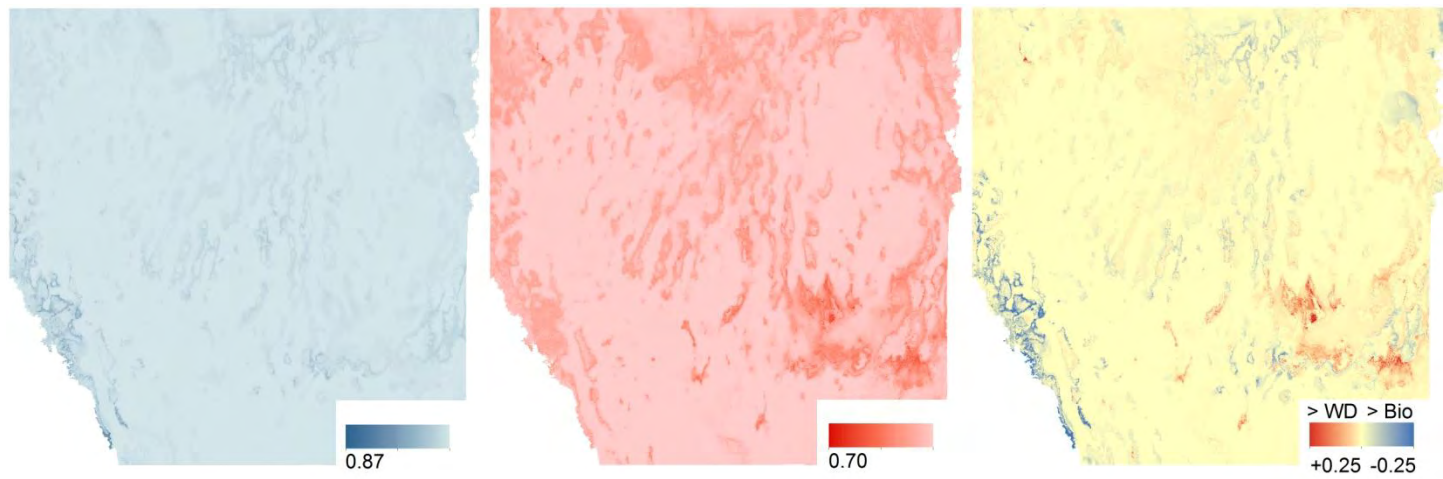
Winterfat (*Krascheninnikovia lanata*)



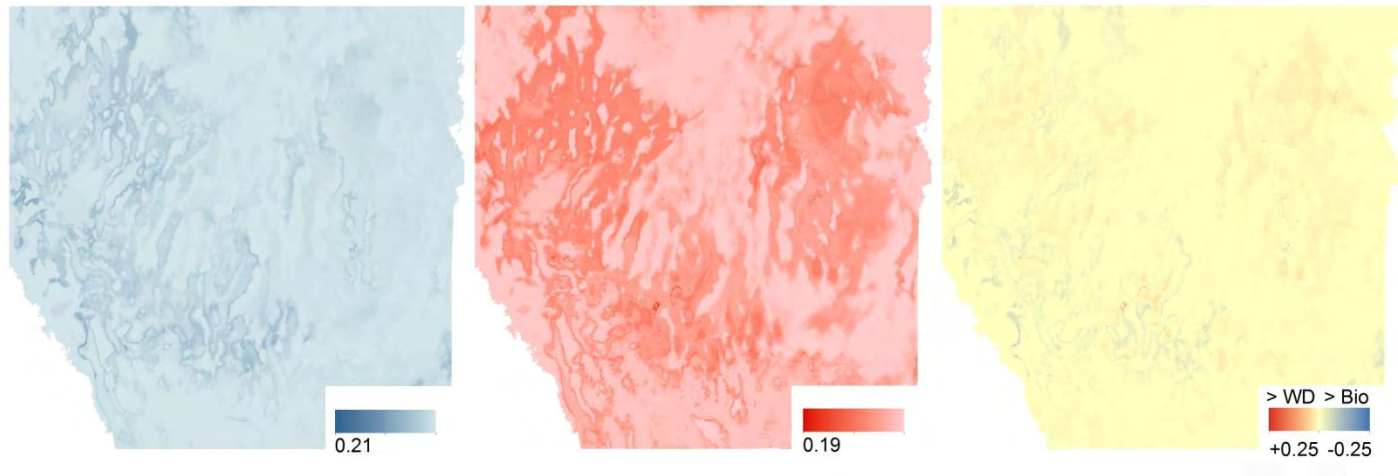
Bailey's greasewood (*Sarcobatus baileyi*)



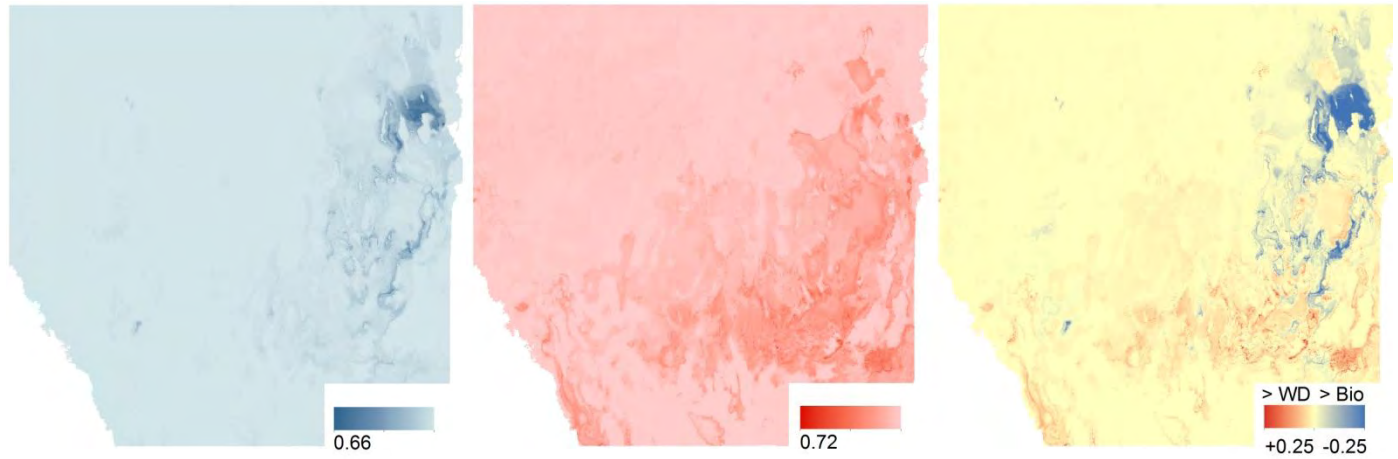
Antelope bitterbrush (*Purshia tridentata*)



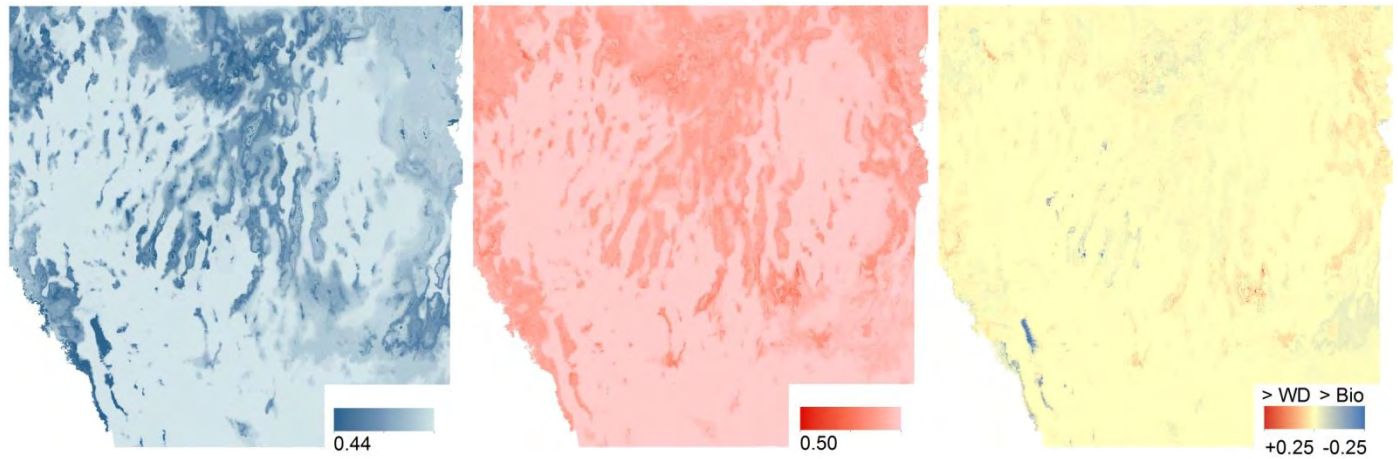
Spiny hopsage (*Grayia spinosa*)



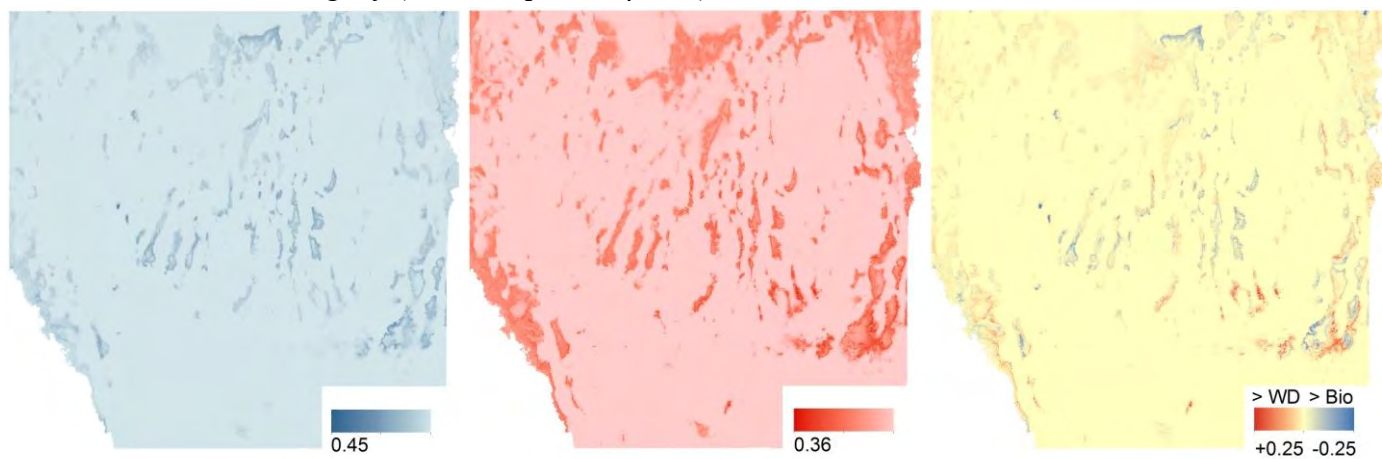
Broom snakeweed (*Gutierrezia sarothrae*)



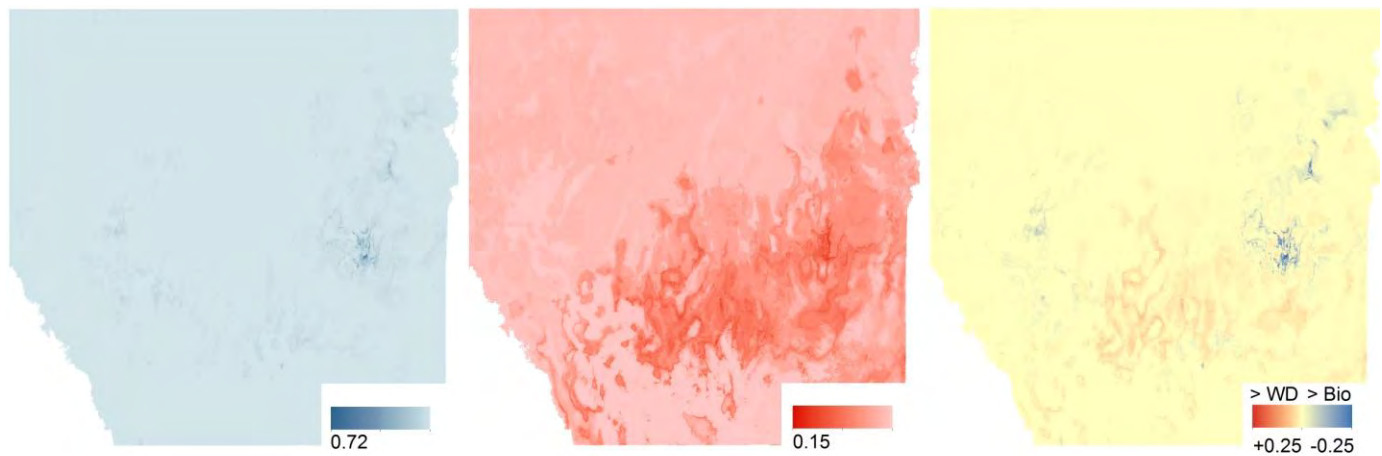
Low sagebrush (*Artemisia arbuscula*)



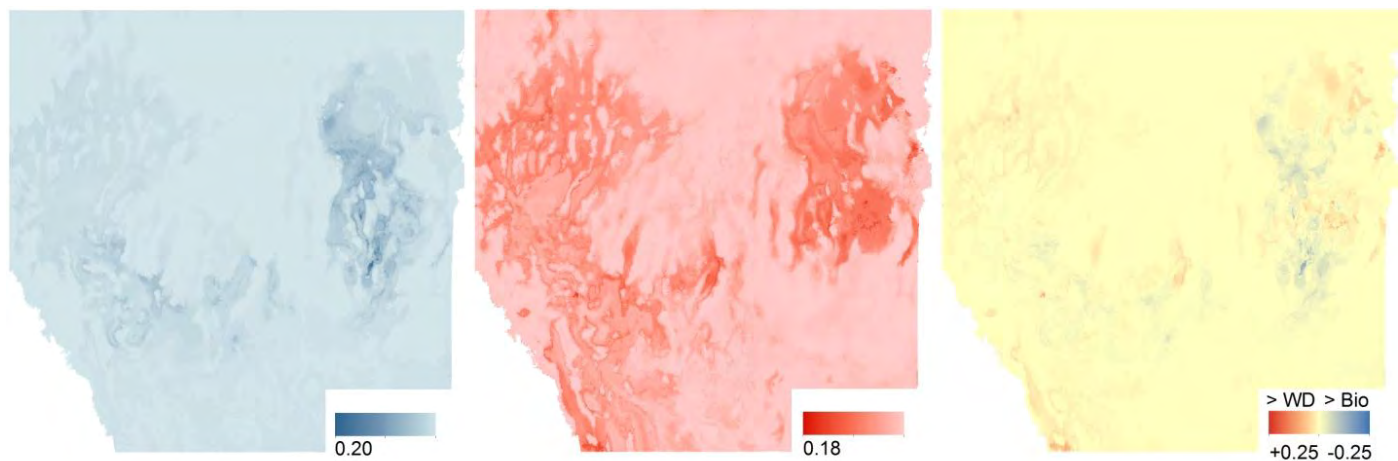
Curl-leaf mountain mahogany (*Cercocarpus ledifolius*)



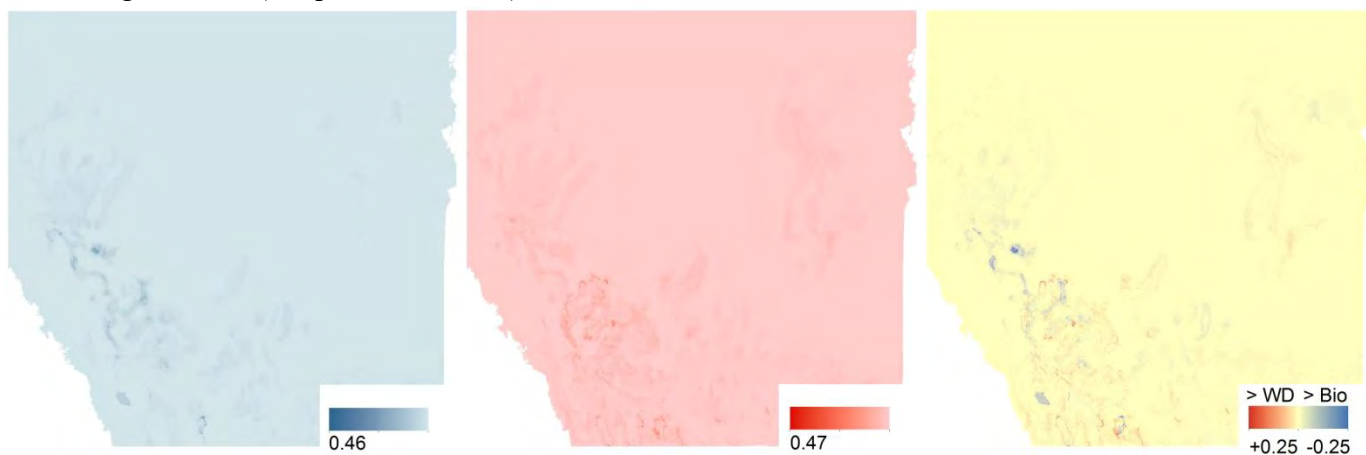
Mormon tea (*Ephedra viridis*)



Littleleaf horsebrush (*Tetradymia glabrata*)



Four-wing saltbush (*Atriplex canescans*)



Species composite mean and standard deviation

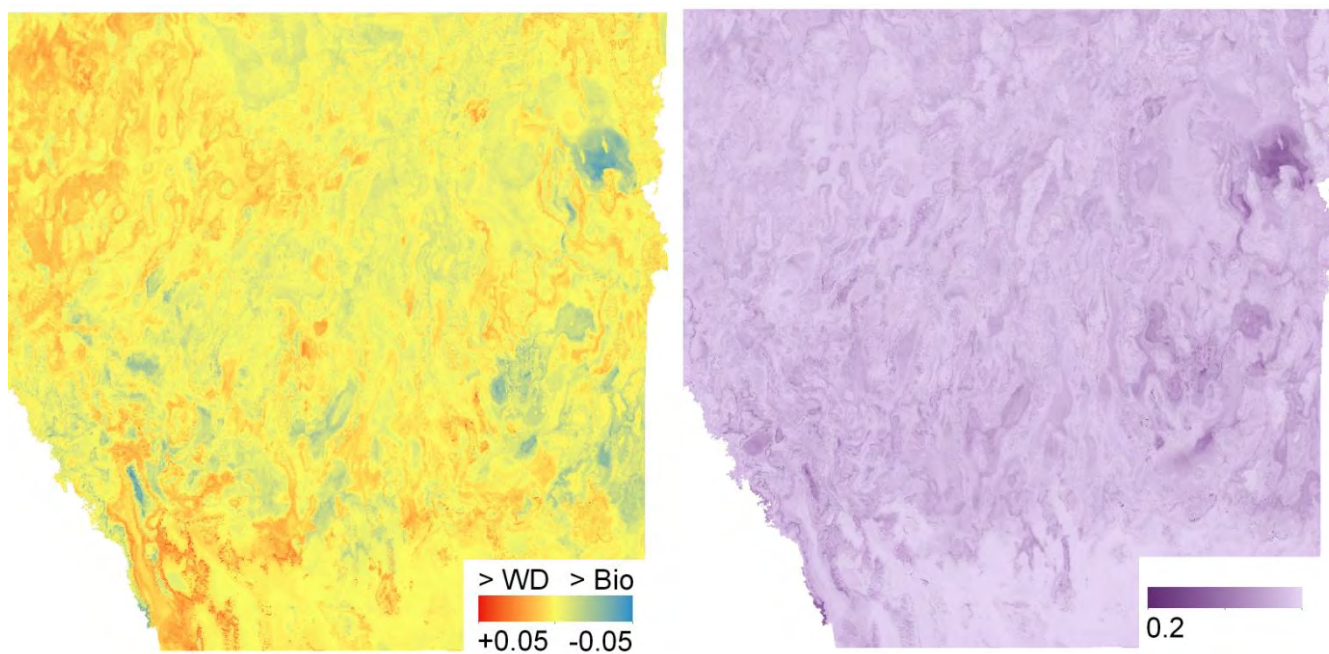


Figure S2 Variable influence for each species distribution model. Values sum to 100 for the bioclimatic model (green) and sum to 100 for the water balance model (orange).

	Temp	Precip	TempRang	PrecipSeas	PrecipDry	CumWCWD	WsAETSpr	FallAET	Ws86c	DecIAET
ARTRW8	33.1	31.5	19.8	6	9.6	50.8	24.7	8.4	5.4	10.6
CHVI8	30.6	15.4	9.7	13.5	30.8	36.8	17.3	14.8	17.4	13.7
ATCO	22.5	12.8	38.6	17.3	8.9	35.6	19.2	15	17.8	12.4
ARNO4	18	24.9	18.7	9.5	28.9	42.4	22.3	12.2	12.8	10.3
ARTRV	22.8	39.6	20.5	4.2	12.9	25.5	57.9	6.1	5.1	5.4
EPNE	33.2	11.1	18.5	19.2	18	33.1	14.2	9.2	30.5	13
SAVE4	20.9	11	46	12.7	9.5	37.3	23.9	21.1	7.7	10
PICRO	22.2	34.8	16	15.3	11.8	26.6	20.8	29.9	12.6	10.1
KRLA2	15	9.8	38.6	22.9	13.7	32.3	12	28.9	11.9	14.9
SABA14	35.4	32.4	12.5	9.5	10.2	32.4	12.6	29.2	11.1	14.7
PUTR2	23.8	20.8	18.6	21.6	15	24.1	26.9	17.2	16.1	15.8
GRSP	21.7	22.1	19.2	19.2	17.8	32.1	22.2	16.2	15.1	14.4
GUSA2	22.2	20.9	17.6	28.6	10.8	18.5	16	17.8	32.6	15.1
ARAR8	24.2	32.8	18	7.8	17.2	22.9	27.6	19	9.9	20.7
CELE3	21.9	36.8	14	9.9	17.4	30.6	37.8	12.9	10.5	8.2
EPVI	14.9	11.4	24.4	40.7	8.5	22.3	17.2	13.6	34.8	12.1
TEGL	24.2	32.8	18	7.8	17.2	32.4	7.7	19.9	21	19.1
ATCA2	16.3	36.4	12.5	22.5	12.3	35	8.7	14.8	18	23.6

Bioclimatic	Water Deficit	
		> 35%
		30-35%
		25-30%
		20-25%
		15-20%
		10-15%
		<10%